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# Middle Carboniferous crustal melting in the Variscan Belt: New insights from U–Th–Pb<sub>tot.</sub> monazite and U–Pb zircon ages of the Montagne Noire Axial Zone (southern French Massif Central)

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## Abstract

In France, the Devonian–Carboniferous Variscan orogeny developed at the expense of continental crust belonging to the northern margin of Gondwana. A Visean–Serpukhovian crustal melting has been recently documented in several massifs. However, in the Montagne Noire of the Variscan French Massif Central, which is the largest area involved in this partial melting episode, the age of migmatization was not clearly settled. Eleven U–Th–Pb<sub>tot.</sub> ages on monazite and three U–Pb ages on associated zircon are reported from migmatites (La Salvétat, Ourtigas), anatectic granitoids (Laouzas, Montalet) and post-migmatitic granites (Anglès, Vialais, Soulié) from the Montagne Noire Axial Zone are presented here for the first time. Migmatization and emplacement of anatectic granitoids took place around 333–326 Ma (Visean) and late granitoids emplaced around 325–318 Ma (Serpukhovian). Inherited zircons and monazite date the orthogneiss source rock of the Late Visean melts between 560 Ma and 480 Ma. In migmatites and anatectic granites, inherited crystals dominate the zircon populations. The migmatitization is the middle crust expression of a pervasive Visean crustal melting event also represented by the “Tufs anthracifères” volcanism in the northern Massif Central. This crustal melting is widespread in the French Variscan belt, though it is restricted to the upper plate of the collision belt. A mantle input appears as a likely mechanism to release the heat necessary to trigger the melting of the Variscan middle crust at a continental scale.

**Keywords:** Migmatites; U–Th–Pb chemical dating; U–Pb geochronology; Variscan Belt; French Massif Central

## 1. Introduction

Crustal melting is a common phenomenon observed in many collision belts, such as Himalaya, Variscan, Appalachian, Qinling–Dabie–Sulu, etc. This feature is often interpreted as the response to thermal relaxation during syn-, or late-collisional unthickening of previously thickened continental crust (e.g. [Le Fort, 1981], [Harris & Massey, 1994], [Faure et al., 2003], [Guillot et al., 2003], [Whitney et al., 2004], [Brown, 2005] and [Faure et al., 2008]).

The Variscan Belt of Western and Central Europe is a Paleozoic collisional orogen that involved large continental masses, such as Gondwana, Baltica and Laurussia, and several

Gondwana-derived microcontinents, such as Avalonia, Armorica, Mid-German Crystalline Rise (e.g. [Paris & Robardet, 1990], [Matte, 1991], [Tait et al., 1997], [Franke, 2000], [Matte, 2001], [Von Raumer et al., 2009], [Johnston & Gutierrez-Alonso, 2010], [Melleton et al., 2010a] and [Melleton et al., 2010b]). The detail paleogeographic and geodynamic reconstructions of the Variscan Belt remain controversial, nevertheless, there is a wide agreement that most of continent amalgamation was completed in the Early Carboniferous. Since the Late Viséan, intracontinental deformation, namely wrenching, thrusting and normal faulting was coeval with intense crustal melting. However, the timing and tectonic setting of these late-collisional events is not clearly settled yet. This paper deals with age constraints on the Middle Carboniferous (Viséan) migmatites and granitoids exposed in the Axial Zone of the Montagne Noire in the southern French Massif Central. New U–Th–Pb<sub>tot.</sub> ages on monazite using Electron Probe Microanalyzer (EPMA) method and U–Pb ages on zircon using High Resolution Secondary Ion Mass Spectrometry (HR-SIMS) were determined to date migmatites, anatectic granitoids and late to post-migmatitic granitic plutons are provided. These data are combined with the available ages of crustal melting known in other Middle Carboniferous migmatitic complexes recognized in the French Massif Central, Massif Armoricain, Vosges and Variscan Pyrénées. A possible geodynamic interpretation of the crustal melting experienced by this segment of the Variscan Belt is discussed.

## **2. An outline of the Variscan French Massif Central**

It is now well accepted that the French Massif Central, which is one of the largest pieces of the Variscan belt of Western Europe, belongs entirely to the northern Gondwana margin. Its bulk architecture results of the stacking of several tectonic–metamorphic units. From bottom to top, and from South to North, five main tectonic–metamorphic units are recognized, namely (Fig. 1): 1) the late Viséan foreland basin; 2) the fold-and-thrust belt; 3) the para-autochthonous unit; 4) the Lower Gneiss Unit; 5) the Upper Gneiss Unit (for detail, see [Faure et al., 2005] and [Faure et al., 2009a] and enclosed references).

The French Massif Central experienced several tectonic–metamorphic events (e.g. [Pin & Peucat, 1986], [Ledru et al., 1989], [Faure et al., 1997], [Faure et al., 2005], [Faure et al., 2009a] and [Faure et al., 2009b]). From Late Silurian to Middle Devonian, the northward continental subduction of the Gondwana margin beneath the Armorica microplate is coeval with the formation of high (or ultra-high) pressure rocks, called D0 event (Lardeaux et al., 2001). Around 380 Ma, the D1 event, referred to as the Eo-Variscan event, is responsible for nappe stacking, exhumation of the high-pressure rocks and the first crustal melting (cf. Faure et al., 2008 and enclosed references). A NE–SW-trending stretching lineation associated with a top-to-the-SW shearing develops during D1 (Roig and Faure, 2000). The stack of nappes is reworked by a D2 event, dated around 360–350 Ma (Late Devonian to Early Carboniferous), and characterized by a top-to-the NW shearing, and coeval with an intermediate pressure–temperature metamorphism ([Roig & Faure, 2000], [Bellot, 2001], [Duguet et al., 2007] and [Melleton et al., 2009]). A Middle Carboniferous thrusting event, called D3, and dated at 340–330 Ma, characterized by South to SW-directed ductile shearing develops only in the southern part of the French Massif Central (e.g. [Arthaud & Matte, 1977], [Ledru et al., 1989] and [Faure et al., 2005]). At the same time, in the northern part of the Massif, the onset of the syn-orogenic extension took place. During the early Late Carboniferous times, around 320–310 Ma, the entire belt experienced extensional tectonics characterized by a pervasive NW–SE striking stretching well recorded by the architecture and emplacement mechanisms of the syn-kinematic granitoids (Faure, 1995).

The French Massif Central exhibits a large amount of magmatic rocks (e. g. [Duthou et al., 1984] and [Pin & Duthou, 1990] and enclosed references). The pre-orogenic Early Paleozoic (mainly Early Cambrian and Early Ordovician) and the Devonian–Early Carboniferous magmatic stages are not presented here (for detail, see Faure et al., 2005 and enclosed references). Since the Middle Carboniferous D3 event, several episodes of crustal melting are recognized. The Late Visean magmatism (ca. 330 Ma), which is represented by the volcanic and sedimentary “Tufs anthracifères” series, is restricted to the northern part of the French Massif Central (Fig. 1). There, rhyolitic to dacitic dykes, microgranites and porphyritic granites represent deeper occurrences of the Late Visean crustal melting event (Faure et al., 2002). During the early Late Carboniferous (Serpukhovian to Baskirian, ca 325 Ma to 310 Ma), crustal melting is represented by a large amount of granitic intrusions. As noticed for a long time (e.g. Didier and Lameyre, 1969), two-mica leucogranites and porphyritic monzogranites are widespread in the northwestern and southeastern parts of the Massif Central, respectively; nevertheless, both types of plutons emplaced in an extensional tectonic setting. Lastly, at the end of the Carboniferous (Moscovian to Gzhelian, 310–300 Ma), cordierite anatectic granite is coeval with the Velay migmatitic event. Furthermore, alkaline acidic and intermediate volcanic rocks emplaced within the intramontane coal basins.

Three generations of migmatites are also recognized during the Paleozoic evolution of the French Massif Central. The oldest one develops from Middle to Late Devonian (ca. 390 to 375 Ma). This Eo-Variscan crustal melting stage is related to the exhumation of the high-pressure metamorphic rocks during the D1 event under ca 600 °C–10 kbar and 800 °C–7 kbar ([Pin & Peucat, 1986] and [Faure et al., 2008]). The second generation of migmatites crops out in the North Cévennes, South Millevaches, and Montagne Noire Axial Zone (Fig. 1). Recent U–Th–Pb<sub>tot</sub> monazite ages in the North Cévennes and South Millevaches range around 330–326 Ma and 337–328 Ma, respectively ([Bé Mézème et al., 2006] and [Faure et al., 2009b]). The third generation of migmatite, dated of Late Carboniferous (ca. 300 Ma) is observed only in the Velay dome (Ledru et al., 2001). Among the Visean migmatites, the Montagne Noire Axial Zone is the largest area. However, there the age of the crustal melting was poorly constrained by the age of the late migmatitic plutons ([Hamet & Allègre, 1976] and [Matte et al., 1998]). In the following, new U–Th–Pb monazite and U–Pb zircon single grain ages for the migmatites, the related anatectic granitoids, and the late migmatitic leucogranites that crop out in the Axial Zone of the Montagne Noire are provided.

### 3. Geological setting of the Montagne Noire

Conversely to the northern part of the French Massif Central, in the southern area, the Devonian tectono-metamorphic events (D0 and D1) are lacking. The Montagne Noire underwent its first deformation during the Middle Carboniferous, from Visean to Serpukhovian (early Namurian). Since (Gèze, 1949) and (Arthaud, 1970), the Montagne Noire is classically subdivided from south to north in southern, axial and northern zones ([Fig. 1] and [Fig. 2]). The Paleozoic series consist of sedimentary rocks ranging from Early Cambrian to Silurian and from early Cambrian to Visean in the northern and southern zones, respectively. These series are involved into kilometre-sized, south verging recumbent folds. The age of the recumbent folding is stratigraphically constrained by the Late Visean to Serpukhovian age of the foreland sedimentary basin in which they emplace ([Engel et al., 1980] and [Feist & Galtier, 1985]).

The Axial Zone consists of metamorphic rocks (micaschists, gneiss, and a small amount of marbles), migmatites, and post-migmatitic plutonic rocks globally arranged in a domal shape,

which is well designed by the foliation attitude and the geometry of the micaschist-gneiss contact ([Gèze, 1949] and [Schuiling, 1960]). To the east, a tight synform of micaschists subdivides the Axial Zone dome into the Espinouze and Caroux sub-domes (Fig. 2). The northeastern limit of the Axial Zone is the Graissessac brittle dextral-normal fault that controls the opening of the Late Carboniferous (Gzhelian) Graissessac coal basin ([Bogdanoff et al., 1984] and [Echtler & Malavieille, 1990]). This brittle structure must not be confused with the Mont-de-Lacaune Fault (MLF) that separates migmatites and granites from high temperature-low pressure micaschist and marbles from south to north, respectively (Fig. 2). The MLF is a ductile structure characterized by a north dipping foliation and a mineral lineation inclined of nearly 45° to the east. Subsequently, the MLF acted as a normal fault with a dextral wrench component. East of Graissessac, Early Permian continental rocks, characterized by red sandstone and mudstone that form the Lodève basin, unconformably overlie the Late Carboniferous terrigenous rocks. The southwestern termination of the axial zone, or Nore massif, is separated from the main part of the dome by the Eocene north-directed reverse fault of Mazamet (Fig. 2). This structure, related to the Pyrenean orogeny, does not alter significantly the Variscan architecture of the Montagne Noire Axial Zone. Although widely acknowledged, the dome structure of the Montagne Noire Axial Zone has been variously interpreted. A discussion of the tectonic setting of this structure is beyond the scope of this paper; the reader can refer to Soula et al. (2001). In spite of the controversy on the Montagne Noire dome emplacement, there is a large consensus on the relative chronology of the tectonic events, namely: 1) recumbent folding and thrusting, 2) crustal melting and doming, 3) post-migmatitic granitoid emplacement, and 4) infill of the Graissessac coal basin. However, the available absolute age constraints are restricted to orthogneiss protoliths and late granitoids ([Hamet & Allègre, 1976], [Ducrot et al., 1979], [Lévêque, 1990], [Matte et al., 1998], [Roger et al., 2004] and [Cocherie et al., 2005b]). Presently, the age of the migmatitic episode is not documented.

#### **4. The Montagne Noire Axial Zone**

Several rock types are recognized in the Montagne Noire Axial Zone. Augen orthogneiss yield zircon U–Pb ages of  $471 \pm 4$  Ma,  $456 \pm 3$  Ma and  $450 \pm 6$  Ma interpreted as the emplacement age of the granites ([Roger et al., 2004] and [Cocherie et al., 2005a]). These Ordovician granitoids intruded a sedimentary sequence of grauwacke, pelite, quartzite, and rare marble with interlayered acidic or mafic magmatic beds ([Demange et al., 1996] and [Alabouvette et al., 2003]). Therefore, the Montagne Noire Axial Zone can no longer be considered as a Precambrian basement covered by a Paleozoic series (outdated interpretation on map by Demange et al., 1996). These rocks experienced a high temperature-low pressure metamorphism with a high temperature gradient of ca 50 °C/km and tight isogrades roughly parallel to the dome shape ([Bard & Rambeloson, 1973], [Demange et al., 1996], [Soula et al., 2001] and [Alabouvette et al., 2003]). This thermal structure agrees well with the progressive crustal melting observed in the centre of the dome. In the dome core, the migmatites progressively grade from metatexites, with leucosomes parallel to the protolith foliation, to diatexite where foliation is absent, and finally to the Laouzas cordierite anatectic granite containing host rocks preserved as metre to kilometre sized septa ([Gèze, 1949] and [Demange et al., 1996]; [Fig. 3] and [Fig. 4]). An earlier, presently undated, high-pressure metamorphism is documented by rare occurrences of kyanite in micaschist, and eclogite blocks in gneiss ([Demange, 1985] and [Alabouvette et al., 2003]).

The northwestern part of the Axial Zone is occupied by the Montalet pluton. This peraluminous granite is represented by two lithologies, a biotite-rich facies and a garnet-rich

one (Fig. 5). The Montalet pluton is considered as an early body ([Demange et al., 1996] and [Alabouvette et al., 2003]). However, as documented in the following section, the magmatic zircon and monazite in this granite yield radiometric ages similar to those obtained for the Laouzas anatectic granite and the migmatites. Several, small sized, late to post-migmatitic plutons intrude the Axial Zone migmatites (Fig. 2). They are either biotite-rich granite (e.g. Anglès granite) or muscovite–biotite  $\pm$  garnet leucogranites (e.g. Soulié, Vialais granites). Lastly, several generations of garnet–tourmaline pegmatite and aplite dykes and sills intrude the above-described metamorphic and plutonic rocks (Collot, 1980).

## 5. Analytical procedures

### 5.1. EPMA dating

Three starting assumptions are required to determine U–Th–Pb<sub>tot</sub> age by using EPMA: 1) common Pb is negligible as compared to the amount of thorogenic and uranogenic lead (Parrish, 1990); 2) no radiogenic Pb loss has occurred since system closure, Pb content is not significantly affected by diffusion (e. g. [Cocherie et al., 1998], [Crowley & Ghent, 1999] and [Zhu & O'Nions, 1999]); 3) a single age is involved at the size level of each individual spot analysis. After comparison with conventional isotopic U–Pb age determinations, it is now accepted that EPMA resolution allows inclusions and altered domains that could potentially contain common Pb to be avoided.

The monazite grains were studied either after extracted by crushing before being mounted in resin or directly by the mean of petrographic thin sections. Both procedures allow the monazite grains to be analyzed using a Cameca SX 50 electron microprobe (cooperated by BRGM and Orléans University, France). Systematic BSE study was performed to investigate monazite micro-texture for all analyzed grains. Obviously, the *in situ* analysis of monazite grains directly in thin sections is more convenient to study age in relation to its petrographic context. This is possible if monazite grains are visible in the thin section and if they are large enough to be sure to avoid border effect when using EPMA.

The analytical procedure for monazite was detailed in (Cocherie et al., 1998) and (Cocherie et al., 2005b); see also Cocherie and Robert, 2008). The interference of YL $\gamma$  on PbM $\alpha$  is subtracted offline by applying a coefficient of interference to the value of Y. The different interference corrections have been validated by dating several monazite samples using this method and using conventional isotopic methods (Cocherie et al., 1998). An accelerating voltage of 20 kV, and a beam current of 100 nA were applied. According to this procedure, the calculated detection limit ( $2\sigma$ ) is 150 ppm for Pb, U and for Th, whereupon an absolute error of 150 ppm is also taken. A systematic relative error of 2% is considered for Th (whose concentration is generally above 7500 ppm), and also for U concentrations above 7500 ppm in order to avoid an unrealistic low error for U-enriched grains (Cocherie and Legendre, 2007). For monazite the standards were galena (PbS) for Pb, uraninite (UO<sub>2</sub>) for U, thorianite (ThO<sub>2</sub>) for Th, endmember synthetic phosphates (XPO<sub>4</sub>) for each rare-earth element (REE) and Y, apatite for P and andradite for Si and Ca.

If two or more homogeneous age domains are separated by a gap lower than the analytical error on each individual spot analysis age, these can be identified by suitable isochron diagrams ([Suzuki & Adachi, 1991], [Cocherie et al., 1998], [Cocherie & Albarède, 2001] and [Suzuki & Kato, 2008]). It has been shown how to provide precise ages  $\pm$  5 to 10 Ma ( $2\sigma$ )

using the most suitable isochron diagram according to the geochemistry of the studied grains (Cocherie et al., 2005b).

A Microsoft Excel add-in program, called *EPMA dating*, was created for determining U–Th–Pb<sub>tot.</sub> ages from electron probe microanalyzer (EPMA) measurements (Pommier et al., 2002). All the parameters needed to calculate mean and intercept ages are computed ready to be plotted using the *ISOPLOT* program (Ludwig, 2003) in order to obtain statistics from suitable diagrams. Finally, *EPMA dating* produces (1) the U–Th–Pb age at the centroid of the best fit line, and (2) the Th–Pb age (intercept with Th/Pb axis) and the U–Pb age (intercept with U/Pb axis) from the Th/Pb vs. U/Pb diagram. All the calculations were done at  $2\sigma$  level. A special care was taken on the MSWD (Mean Squared Weighted Deviation) which must be below  $1+2/(2/f)^{0.5}$  ( $f$ : degree of freedom = number of analyses – number of dated events) in order to validate age calculation (Wendt and Carl, 1991) for a single age population. Using such a diagram, a single population must also lead to a regression line showing the same slope as the corresponding theoretical isochron. In other words, the U–Pb and Th–Pb age intercepts must be similar within the uncertainties on these ages.

## 5.2. SIMS dating

Zircons were dated using the SIMS (IMS 1270) at CNRS, CRPG Nancy (France), according to the procedure described by (Deloule et al., 2001) and (Deloule et al., 2002). The areas analyzed ( $\sim 25 \mu\text{m}$ ) were selected after studying images of the grains obtained by cathodoluminescence and by transmitted light microphotography. The geostandard 91500 zircon from Ontario (Canada), dated at 1065 Ma (Wiedenbeck et al., 1995) is used for  $^{238}\text{U}/^{206}\text{Pb}$  calibration. In general, areas very rich in U ( $> 2000\text{--}3000 \text{ ppm}$ ) were not selected in order to avoid, firstly, moving away from the area of validity of the U–Pb calibration line and, secondly, the risk of losses of radiogenic Pb, in relation to the metamictisation. For the relatively recent zircons ( $< 1000 \text{ Ma}$ ), the imprecision of the  $^{206}\text{Pb}/^{204}\text{Pb}$  ratio becomes critical; one then uses the Concordia diagram (Tera and Wasserburg, 1972), modified by Compston et al. (1992), in which one plots the  $^{207}\text{Pb}/^{206}\text{Pb}$  and  $^{238}\text{U}/^{206}\text{Pb}$  ratios, not corrected for common Pb. In the absence of common Pb, the analyses of areas not affected by recent Pb\* loss after the crystallisation of the zircon or by inherited cores are spread along this Concordia (Williams, 1998). Although the variable quantities of common Pb adversely affect the values of the two ratios, these kinds of points form a straight line passing through the composition of common Pb ( $^{207}\text{Pb}/^{206}\text{Pb}$ ) at the estimated age of the system given, in a first approximation, by the average  $^{238}\text{U}/^{206}\text{Pb}$  ages stemming from the concordant analyses. The extrapolation of this line on the Concordia defined the sought age. This is what is called a correction of common Pb by the  $^{207}\text{Pb}$  method and not by the  $^{204}\text{Pb}$  method, as in the case of the conventional diagram. Using this correction method, one calculates the  $^{206}\text{Pb}^*/^{238}\text{U}$  ( $\text{Pb}^* = \text{radiogenic Pb}$ ) ratios for each point. Whatever the used analytical approach, all the calculations were made at  $2\sigma$  (95% confidence limit) using the *ISOPLOT* program (version 3) of Ludwig (2003). On the other hand, in the case of the data obtained using the SIMS, the uncertainties are given at  $1\sigma$  in the corresponding table and, in the same way, the error ellipses are given at  $1\sigma$  in order to make the figures easier to read.

## 6. Geochronological results

New geochronological data are reported on coexisting monazite and zircon, for migmatites, anatectic granitoids and late-migmatitic plutons from the Axial Zone of the Montagne Noire. Monazite and zircon crystals were selected from a mineral separate, after crushing the

samples, mounted in epoxy mount, and polished to approximately half thickness. For five of the samples, monazite was analyzed in thin section in its petrographic context. Monazites were imaged with a backscattered electron detector (BSE) in a scanning electron microscope (SEM) before collecting U–Th–Pb<sub>tot</sub> data with an EPMA. Zircons were imaged with a cathode-luminescence detector (CL) in a SEM before U–Pb isotopic analysis by SIMS.

## 6.1. Migmatites

### 6.1.1. *La Salvetat migmatite*

#### 6.1.1.1. *U-Th-Pb<sub>tot</sub> monazite dating*

This rock is a diatexite (Fig. 4) sampled near the La Salvetat village (GPS location: N43.61754, E 2.75685). Five monazite grains were analyzed on two thin sections. The grains, with sizes ranging from 50 to 150  $\mu\text{m}$ , are included in biotite (Fig. 6a, b). Their shapes are subhedral or rounded, and rarely cracked. The back-scattered electron (BSE) scan electron microscope (SEM) images of the analyzed grains show a well developed patchy zoning structure but no regular growth zoning (Fig. 6c).

The BSE images are representative of the electronic density, thus Th- and U-enriched domains appear clearer than other domains; note that generally the weight percentage of Y is too low to provide significant change in BSE images. Th- and U-enriched domains can show similar Th/U ratio as those observed in Th- and U-depleted domains, leading to contrasting BSE images associated to constant Th/U ratio. Thus, rather constant Th/U ratio (Fig. 7; Table 1) is observed for La Salvetat monazite despite BSE images show contrasting domains. This constant Th/U ratio results in a large error envelop around the regression line. Despite the lack of a large spread of the representative spot analyses, the slope of the regression line is very close to the theoretical isochron. This demonstrates that all the spot analyses are similar in age within the analytical uncertainties. A mean age of  $327 \pm 7$  Ma was calculated at the centroid of the population. The MSWD (1.5) is only slightly above the upper limit when a population of 48 analyses is taken into account (1.4, see Wendt and Carl, 1991), because none of the data were rejected. Therefore the analyzed grains will be considered as a single population. A detailed examination of the data and Th/Pb vs. U/Pb plot makes visible three groups of composition according to the Th/U ratio. However, this diagram demonstrates that they cannot be distinguished in age (within the analytical resolution of EPMA dating) because the regression line has the same slope as the theoretical isochron at 327 Ma. This age is interpreted as the crystallization age of the monazite. No inherited domain was detected.

#### 6.1.1.2. *U-Pb zircon dating*

This sample contains abundant oscillatory zoned elongate zircon crystals. Most of the 34 SIMS analyses performed in 26 crystals plot along or near the Concordia curve, and define a cluster at around 550 Ma. One discordant analysis yield a  $^{207}\text{Pb}^*/^{206}\text{Pb}^*$  age of 2365 Ma. Special attention has been paid to analyze the youngest possible domains of the zircons, including the pyramidal extremities of the crystals. In spite of this effort, only three analyses seem to date the Variscan migmatization. These three analyses project on the Concordia curve at an age of  $306 \pm 6$  Ma ( $^{207}\text{Pb}$  corrected  $^{206}\text{Pb}/^{238}\text{U}$  age).



One out of three is significantly discordant due to common Pb contribution (not corrected in this diagram), but it can be considered in the calculation assuming a mixing line with the common Pb composition in the Variscan times. The calculated age is  $306 \pm 6$  Ma. The cluster of analyses around 550 Ma can be interpreted in two ways. (1) A single age was recorded and various degree of radiogenic Pb\* loss results in a shift of more than 50% of the analyses. (2) Two main inherited protoliths were recorded at  $562 \pm 6$  Ma (MSWD = 0.8,  $n = 11$ ) and  $530 \pm 5$  Ma (MSWD = 1.4,  $n = 13$ ). A third one could also be dated at  $478 \pm 13$  Ma. We think that an extended recent Pb\* loss should provide a wider range of variation of the data along the  $^{238}\text{U}/^{206}\text{Pb}$  axis. Therefore, in addition to the 562 Ma age, 530 and 478 Ma could also be significant on a geological point of view (see discussion section, below).

It must be underlined that only three spot analyses are considered for the age calculation at 306 Ma. In addition, only two of them are concordant at  $304 \pm 8$  and  $315 \pm 16$  Ma ( $2\sigma$ ) while monazite was dated at  $327 \pm 7$  Ma ( $2\sigma$ ). Thus, within the analytical uncertainties and considering only these two spot analyses on zircon overgrowths we cannot really conclude that monazite grains crystallized significantly before or at the time of zircon recrystallization. The usual conception is that zircon ages tend to be older than monazite ages within the same rock. Furthermore, recent experimental data (Kelsey et al., 2008) led to the same conclusion. However, we also studied numerous monazite and zircon samples from natural migmatites. It was demonstrated that monazite can remain as a residual accessory phase while inherited zircon grains were dated at the same age but also show overgrowth during the migmatitic stage, 80 Ma later, not visible on monazite grains (Cocherie et al., 2005b). Several unpublished data from leucocratic granites show the same trend. It seems to demonstrate first that monazite is a very robust mineral up to the temperature of zircon overgrowth and second that there is a lack of some chemical elements required for the overgrowth of preexisting monazite during migmatization.

### 6.1.2. Ourtigas migmatite

#### 6.1.2.1. U-Th-Pb<sub>tot</sub> monazite dating

Two samples (Ourt1 and Ourt2) of the metatexite developed at the expense of an augen orthogneiss, the foliation of which remains well preserved (Fig. 3), were collected at the Ourtigas pass (GPS location: N43.63657, E2.97628). The monazite grains are quite similar to those from the Salvetat migmatites (Fig. 8). Seven and three monazite grains from Ourt1 and Ourt2, respectively were separated and mounted in polished sections. Two age populations associated with different monazite grains can be identified from the Ourt1 sample. The first population consists of 29 analyses obtained on three grains, which define a consistent regression line in the Th–Pb vs. U–Pb diagram (Fig. 9). The intercept ages are rather well defined (U–Pb age:  $314 + 14\text{--}16$  Ma and Th–Pb age:  $347 + 26\text{--}22$  Ma) and similar within the errors. A mean age of  $326 \pm 4$  Ma (MSWD = 0.71) was calculated at the centroid of the population. Four grains define the second population that provide a distinct regression line when plotting 30 analyses in the Th–Pb vs. U–Pb diagram (Fig. 9). The intercept ages are similar within the errors on these ages: U–Pb age:  $420 + 48\text{--}51$  Ma and Th–Pb age:  $468 + 17\text{--}15$  Ma. A mean age of  $458 \pm 9$  Ma (MSWD = 0.62) was calculated at the centroid of the population. No textural feature allows us to distinguish the two monazite populations yielding 326 Ma and 458 Ma ages. In addition, it is surprising that no overgrowths were observed on the cores of old monazite grains. Such overgrowths were observed for instance on monazite grains from French Massif central in Puylaurent migmatite (Bé Mézème et al., 2006) or in St Laurent granite (Cocherie et al., 2005b). We suggest that the older grains were not in

equilibrium with the magma from which the younger grains crystallized. They can be inherited from the wall rocks during the emplacement the migmatitic melt during Variscan times.

Fifty-four analyses were recorded from three monazite grains from Ourt2 sample (Fig. 7). The Th/U ratio changes significantly, leading to a well-defined slope of the regression line. This slope is very similar to that of the theoretical isochron. A mean age of  $524 \pm 7$  Ma (MSWD = 1.3) was calculated at the centroid of the population.

## 6.2. Anatectic granitoids

### 6.2.1. *Laouzas cordierite granite*

#### 6.2.1.1. *U-Th-Pb<sub>tot</sub> monazite dating*

The Laouzas cordierite granite represents one end-member of the progressive evolution from partially melted country gneiss to anatectic granite through metatexite and diatexite. The dated sample (GPS location: N43.63363 E2.75685) contains numerous enclaves of the metamorphic country rocks (Fig. 4). Alike for the migmatite, in the Laouzas granite, the monazite occurs as inclusion in biotite. Euhedral monazite grains with size ranging from 50 to 150  $\mu\text{m}$ , commonly exhibit a patchy zoning, and more rarely contain thorite inclusions. The monazite grains of the Laouzas cordierite granite were directly investigated on thin section. Fifty-four spot analyses were performed on five grains by using the EPMA technique. Due to rather constant Th/U ratio, the spot analyses are clustered in the Th/Pb vs. U/Pb diagram (Fig. 10) leading to large errors on the intercept ages (U–Pb age:  $321 + 65 - 74$  Ma and Th–Pb age:  $336 + 23 - 20$  Ma). However, it is noteworthy that the slope of the regression line is very close to that of the theoretical isochron. A mean age of  $333 \pm 6$  Ma (MSWD = 0.65) was calculated at the centroid of the population ( $n = 54$ ).

#### 6.2.1.2. *U-Pb zircon dating*

Zircons from this sample are elongate and oscillatory zoned (Fig. 11b). Thirty-two SIMS analyses were performed on 27 crystals (Table 2c; Fig. 12b). Most of the analyses plot around 550–520 Ma. In spite of efforts to retrieve a Variscan intrusion age, only two analyses can be related to crystallization of the granite itself. The age of  $299 \pm 8$  Ma derived from these two analyses is probably affected by Pb\* loss and is not regarded as a reliable estimate. Taking into account the uncertainties on the oldest spot analyses, more than one age must be considered. Again, we favour the hypothesis considering that two main inherited protoliths were recorded at  $552 \pm 5$  Ma (MSWD = 1.4,  $n = 18$ ) and  $520 \pm 7$  Ma (MSWD = 2.0,  $n = 12$ ). However, similarly to La Salvetat zircons, because the second group of analyses is not concordant, we must consider the hypothesis that some radiogenic Pb\* loss happened, making the age of 520 Ma not significant. Nevertheless, as for La Salvetat inherited zircons, a recent Pb\* loss should provide a much wider range of variation of the data along the  $^{238}\text{U}/^{206}\text{Pb}$  axis. Therefore, the 552 Ma and 520 Ma ages can be both significant from a geological point of view.

### 6.2.2. *Montalet granite*

Three samples of the Montalet granite have been analyzed. Two of them belong to the biotite-rich facies and one to the garnet-rich facies (Fig. 5). In all three samples, monazite occurs as

inclusion in biotite or more rarely as an interstitial phase along biotite and K-feldspar grain boundaries. The monazite grains of the garnet-rich and biotite-rich facies are anhedral and without zoning (Fig. 13) which does not mean necessarily that they are not chemically zoned. Indeed, the large scatter of points in the Th/Pb vs. U/Pb diagram for one the Montalet biotite-rich facies and the garnet-rich facies suggest that the monazite is chemically zoned. They have been studied directly on thin sections. Furthermore, the garnet-rich facies has also been dated on separated zircon by SIMS.

#### **6.2.2.1. Garnet rich-facies**

This sample (GPS location: N43.61897, E2.60599) is an oligoclase, K-feldspar, fine-grained leucocratic monzonite with millimetre-sized almandine (Fig. 5). Fifty-eight spot analyses were performed on five grains by using the EPMA technique. Due to a rather small amount of Th and U, the ellipse errors associated to each individual analysis are large (Fig. 10). The theoretical isochron remains within the uncertainty defined by the error envelop of the regression line. A mean age of  $327 \pm 7$  Ma (MSWD = 0.95) was calculated at the centroid of the population.

The zircon grains from the same sample (garnet rich facies) were also investigated (Fig. 12c; Table 2b). Analyzed grains are generally elongated (100 to 150  $\mu\text{m}$ ) needles but they exhibit contrasting internal structure. The cathodoluminescence images of representative grains are shown on Fig. 11c. Some euhedral or subhedral grains exhibit a core and rim structure.

Twenty-five spots were performed on 21 zircon grains. A significant number of analyses show common-Pb contribution in the uncorrected Tera and Wasserburg diagram (Fig. 12c). However, a nice mixing trend between common Pb and radiogenic Pb provide an intercept age at  $324 \pm 3$  Ma (MSWD = 1.4) using 9 analyses. This age is similar to the age calculated from the monazite of the same granite. Abundant inherited crystals yield dates at around 460 Ma, 1000 Ma and even 2700 Ma (Table 3).

#### **6.2.2.2. Biotite-rich facies**

The dated sample (GPS location: N43.62194, E2.50402) is an oligoclase, biotite-rich medium-grained monzogranite with abundant K-feldspar megacrysts ([Fig. 5] and [Fig. 13]). Monazite has been investigated directly on thin section. The grains are rather small (around 40–70  $\mu\text{m}$ ). Some of them show a patchy zoning but no inherited cores, sixteen grains have been analyzed. The U/Pb and Th/Pb intercept ages are similar within errors:  $333 \pm 67/-55$  Ma and,  $326 \pm 6.9/-6.6$  Ma respectively (Fig. 10). In addition, the MSWD of 1.1 indicates that the population of 159 analyses is statistically homogeneous within the analytical errors (Wendt and Carl, 1991). Thus, the monazite grains lead to a mean age of  $333 \pm 4$  Ma (at 95% confidence level) calculated at the centroid.

The second sample is also a medium-grained mesocratic porphyritic granodiorite petrologically similar to the previous one (GPS location: N43.64770, E 2.59933). The chemical composition of the monazite grains was very favourable for U–Th–Pb<sub>tot</sub> geochronology since they exhibit a high U and Th content, and a wide range of variation for the U/Th ratio. Consequently, the ellipses errors are small and the data plot favourably in the Th/Pb vs. U/Pb diagram (Fig. 10) showing a narrow error envelop. The intercept ages are very well defined (U–Pb age:  $482 \pm 21-23$  Ma and Th–Pb age:  $507 \pm 13-12$  Ma) and similar

within the errors. A mean age of  $499 \pm 6$  Ma (MSWD = 1.1) was calculated at the centroid of the population ( $n = 96$ ).

### **6.3. Late to post migmatitic granites**

Among the numerous kilometre-sized plutons that intrude into the migmatites, the Anglès, Vialais, and Soulié massifs have been dated by monazite U–Th–Pb monazite method.

#### **6.3.1. Anglès pluton**

The analyzed sample (GPS location: N 43.56441, E2.55789) is a medium grained monzonite in which biotite exhibits a clear planar preferred orientation. Three separated monazite grains were mounted in polished section. The SEM back-scattered electron images show subeuhedral, cracked grains without chemical zoning (Fig. 13c and d). A well-defined regression line that perfectly fits with the theoretical isochron was obtained after plotting 56 analyses in the Th–Pb vs. U–Pb diagram (Fig. 15a). The intercept ages are rather well defined (U–Pb age:  $322 + 34\text{--}40$  Ma and Th–Pb age:  $327 + 31\text{--}26$  Ma) and they are identical within the errors. In other words the regression line and the theoretical isochron are almost identical. A mean age of  $325 \pm 7$  Ma (MSWD = 1.2) was calculated at the centroid of the population.

#### **6.3.2. Vialais pluton**

This medium-grained leucogranite (GPS location: N 43.64436 E 2.96685) contains rare garnet grains and does not exhibit a mineral preferred orientation. Five separated monazite grains were mounted in polished section. Th and U contents are high and the Th/U ratio show a wide range of variation that makes the Th–Pb vs. U–Pb representation very precise (Fig. 15b). The regression line and the theoretical isochron are very close. A mean age of  $320 \pm 3$  Ma (MSWD = 0.92) was calculated at the centroid of the population.

#### **6.3.3. Soulié pluton**

The analyzed sample (GPS location N43.53352 E2.66002) is a two-micas monzogranite with globular quartz and large muscovite grains. The SEM BSE images reveal subeuhedral, slightly fractured monazite grains without chemical zoning (Fig. 13c). Three separated monazite grains were mounted in polished sections. A well-defined regression line was obtained after plotting 68 analyses in the Th–Pb vs. U–Pb diagram (Fig. 15). The intercept ages are rather well defined (U–Pb age:  $298 + 22\text{--}23$  Ma and Th–Pb age:  $325 + 9\text{--}8$  Ma) and they are identical within the errors. The regression line and the theoretical isochron are almost identical. A mean age of  $318 \pm 4$  Ma (MSWD = 1.08) was calculated at the centroid of the population.

## **7. Discussion**

### **7.1. Geological significance of the new results**

The U–Th–Pb<sub>tot</sub> monazite ages presented above are consistent and geologically meaningful. The ages of the La Salvetat, and Ourtigas migmatites around 327–326 Ma are in agreement with that of the Laouzas anatectic granite at  $333 \pm 6$  Ma. These late Viséan ages are also similar to the emplacement age of the Montalet pluton dated on monazite at  $327 \pm 7$  Ma and

on zircon at  $324 \pm 3$  Ma. Therefore, data available today do not support the assumption that the Montalet granite is an early pluton ([Demange et al., 1996] and [Alabouvette et al., 2003]). Furthermore, the coeval character of the migmatites and anatectic granitoids is consistent with petrofabric and AMS analyzes showing that the Montalet granite, Laouzas granite, and migmatites experienced the same sub-solidus to post-solidus deformation stages during the doming of the Montagne Noire Axial Zone (Charles et al., 2009). The new results demonstrate the abundance of Neoproterozoic, Cambrian and Ordovician ages inherited zircon and monazite grains. Such “old” ages are already identified in the Montagne Noire, for instance in the Héric, Plaisance, Mazamet and Mendic orthogneiss (Fig. 16, [Hamet & Allègre, 1976], [Ducrot et al., 1979], [Roger et al., 2004] and [Cocherie et al., 2005a]). The ca. 560–530 Ma ages are interpreted as the crystallization ages of the granite protolith of the migmatites and anatectic granites. Furthermore, the Early Ordovician ages (ca 480 Ma) also correspond to the alkaline granites related to the rifting of Armorica from Gondwana, and transformed into orthogneiss during the eo-Variscan D0 and D1 events. Moreover, Paleoproterozoic and Archean zircons, recovered from the migmatites and anatectic granitoids, represent unmelted relict grains inherited from the source of the Early Paleozoic magmatic rocks. Archean or Paleoproterozoic ages are now well identified in many Variscan French massifs, such as the Pyrénées or Massif Central ([Deloule et al., 2002], [Cocherie et al., 2005a], [Melleton et al., 2010a] and [Melleton et al., 2010b]).

Our ages of the late migmatitic granites range from  $325 \pm 7$  Ma to  $318 \pm 4$  Ma. These dates are also in agreement with field and petrofabric observations showing that these plutons emplaced within an already foliated migmatite. The monazite age at  $325 \pm 7$  Ma of the Anglès biotite massif is close to the ages of the migmatite and contemporaneous anatectic granites. In addition to field observations showing numerous biotite enclaves, host rock xenoliths and poorly defined boundaries, this date is in agreement with a late-migmatitic setting for the Anglès pluton.

The Soulié biotite–muscovite granite does not present a macroscopic mineral preferred orientation. Its chemistry and young age at 318 Ma suggest that this massif belongs to the most differentiated magmas such as the late pegmatite dykes that emplaced during the late magmatic stages in the Axial Zone of the Montagne Noire.

Lastly, the  $320 \pm 3$  Ma monazite age presented in this paper for the Vialais pluton is slightly younger than the zircon U–Pb TIMS age at  $327 \pm 5$  Ma obtained previously for the same massif (Matte et al., 1998). However, these two ages are identical within the analytical uncertainties. As presented above, available maps and field investigations clearly show that the Vialais pluton intrudes at high angle into the foliated migmatites (Bogdanoff et al., 1984).

## **7.2. Late Visean partial melting in the French Variscan belt**

The Montagne Noire Axial Zone represents the largest area of the Visean migmatites in the French Massif Central. In the recent years, several migmatite occurrences have been radiometrically dated of Late Visean in several parts of the French Massif Central (Fig. 1). In northern Cévennes, the migmatites that crop out along the southern edge of the Late Carboniferous Velay dome have often not been distinguished from the ca 300 Ma Velay migmatites (Ledru et al., 2001). However, several lines of evidence argue for an earlier crustal melting event (Montel et al., 1992). The monazites from the northern Cévennes migmatites yield U–Th–Pb ages at ca 330–325 Ma, and they are intruded by two-micas granites dated at 320–315 Ma ([Cocherie et al., 2005b] and [Bé Mézème et al., 2006]). To the west, migmatites

that crop out in the southern part of the Millevaches massif have been dated between 337 and 328 Ma by the U–Th–Pb method on monazite (Faure et al., 2009b). Lastly, in the northeastern French Massif Central, the Autun migmatites (Fig. 1) yield monazite chemical ages at ca 330 Ma (Choulet, personal communication). Zircon from migmatites found in the central part of the Vosges yields U–Pb ages at ca 340 Ma (Schaltegger et al., 1999; Fig. 17).

More to the South, the Variscan Pyrenees presents great similarities with the Montagne Noire area. There, migmatitic domes and anatectic granitoids are recognized for a long time in the Hospitalet, Aston and Canigou massifs but smaller outcrops are also identified (e. g. [Zwart, 1986], [Soula et al., 1986], [Vissers, 1992], [Carreras & Capella, 1994], [Mezger, 2005] and [Denèle et al., 2007]; Fig. 17). Although the age of the migmatite is not available yet, the upper limit for crustal melting is provided by the Late Carboniferous age of the granitoids intruding the migmatite (Soula et al., 1986).

In the southern part of the Massif Armoricain, which also belongs to the northern Gondwana domain, migmatites are well exposed in the Vendée and Morbihan areas (Fig. 17; [Cogné, 1960], [Audren, 1990], [Goujou, 1992] and [Brown, 2005]). The Morbihan migmatites have recently been dated at ca 330–325 Ma by EPMA method on monazite (Turrillot et al., 2009). Furthermore, the migmatites that form the main part of the Léon domain in the northwest Armorican Massif yield also 335–327 Ma monazite chemical ages (Faure et al., 2010).

East of the Cévennes fault, Variscan migmatites are also well exposed in the external crystalline massifs of the Western Alps, and in the Maures, and Corsica–Sardinia Massifs ([Guillot & Ménot, 2009], [Corsini & Rolland, 2009] and [Rossi et al., 2009]; Fig. 17). In all of these areas, radiometric constraints on the migmatization are still rare. Recently, monazite chemical dating of the migmatites that crop out in the Maures Massif indicates a Late Visean–Serpukhovian age (Melleton et al., in review).

In addition to the migmatitic domains, the Late Visean partial melting is well represented in the Variscan French massifs. Volcanic and volcano-sedimentary rocks of the “Tufs anthracifères” series crop out widely in the northern part of the French Massif Central and Vosges ([Fig. 1] and [Fig. 17]). Similar rocks are also developed in the Châteaulin and Laval basins in the North-Central Armorican domain (Fig. 17). Therefore, at the scale of the entire French Variscan belt, the Late Visean crustal melting corresponds to an important episode responsible for the emplacement of huge volumes of melts. Fig. 17 shows that the distribution of the areas that experienced the Late Visean crustal melting are widely scattered.

In most of areas, the tectonic setting of the Late Visean migmatites is poorly constrained or remains controversial. Dome structures are documented in a few places such as the Montagne Noire, the Pyrénées or in Vendée. Several tectonic interpretations, namely compressional, extensional, trans-tensional or diapiric settings have been proposed, a discussion of these models is beyond the scope of this paper (cf. [Van Den Driessche & Brun, 1992] and [Soula et al., 2001] for extensive reviews).

### **7.3. A mantle origin for the heat?**

The Late Visean partial melting can be situated in a larger perspective at the scale of the French Variscan segment. Lava flows, and migmatitic–plutonic complexes represent the

shallow and deep records of the same crustal melting event. Tectonically, melting developed during the North-directed and South-directed thrusts observed in the Ardenne and Montagne Noire-Pyrénées domains, respectively, whereas, in the northern part of the French Massif Central, the Late Visean Tufts anthraciferous magmatism is coeval with the onset of the NW–SE post-collisional extension (e.g. Faure et al., 2005). The Late Visean crustal melting occurred after the main phase of continental collision between three continental masses, namely from North to South, Laurussia, Mid-German Crystalline Rise-Léon, and Gondwana. The Rheic-North Phyllite Zone suture and Le Conquet–Penzé–Tepla suture separate these continents ([Fig. 16] and [Fig. 17]). As a whole, the Carboniferous tectonic evolution is controlled by a south-directed subduction (e.g. [Matte, 1991], [Matte, 2001], [Franke, 2000], [Faure et al., 2002], [Faure et al., 2009a] and [Faure et al., 2009b]). In this framework, the migmatitic, plutonic and volcanic complexes presented above are located in the upper plate, i.e. in the units derived from Gondwana, of the convergent system. Thus, as sometimes emphasized, the Carboniferous Variscan geodynamic setting can be compared to an Andean orogen (e.g. Schulman et al., 2009).

The question of the heat source able to produce such a continental-scale melting arises. A first explanation would link the heat release with the un-thickening phase that generally follows continental collision. However, another possibility is to consider a mantle heat contribution. This explanation requires the thinning of the subcontinental mantle lithosphere and its replacement by a hot asthenospheric mantle that might be achieved by various mechanisms such as delamination, magmatic underplating or convective ablation.

Great attention has been paid to the Andean Belt as a typical example of an orogen developed in the upper plate of a converging system. Numerous works emphasized the importance of partial melting in the thick crust of the Andes Cordillera (e.g. [ANCORP Working Group, 2003], [Schilling et al., 2006] and [Oncken et al., 2006] and enclosed references). Interestingly, numerical modelling showed that heat necessary to produce molten layers in the crust must be provided by the asthenosphere mantle, and that convection of partially molten crust must have played an important role to account for the present magmatism of the Central Andes (Babeyko et al., 2002). Tentatively, we speculate here that a similar mechanism might be responsible for the development of the widespread crustal melting in the Variscan belt. After the main continental amalgamation was completed in Early Carboniferous, the lithosphere mantle was thinned and heat transfer from the asthenosphere triggered the melting of the middle crust giving rise to intracrustal convection and development of anatectic granite–migmatitic complexes throughout the entire belt (Fig. 18).

## 8. Conclusion

This EMPA and SIMS geochronological study on monazite and zircon, respectively allows us to clarify the age of the partial melting in the Montagne Noire Axial Zone. The Visean (ca 333–326 Ma) age of the migmatites and anatectic granitoids is in good agreement with the ages recovered in other massifs of the French Variscan Belt. An Andean-type orogeny, in which the migmatitic regions belong to the upper plate appears as a possible model to account for the geodynamic setting of this Carboniferous crustal melting event. However, the mechanisms of mantle thinning remain a matter of debate. Furthermore, zircon yields SIMS Archean, Paleoproterozoic, Neoproterozoic–Cambrian and Ordovician ages. These grains are interpreted as inherited from the protolith sources of the melts. They suggest that the Axial Zone of the Montagne Noire contains numerous Precambrian and Early Paleozoic plutonic rocks. The tectonic setting of these protoliths remains to be clarified.

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## **Figures and Tables**



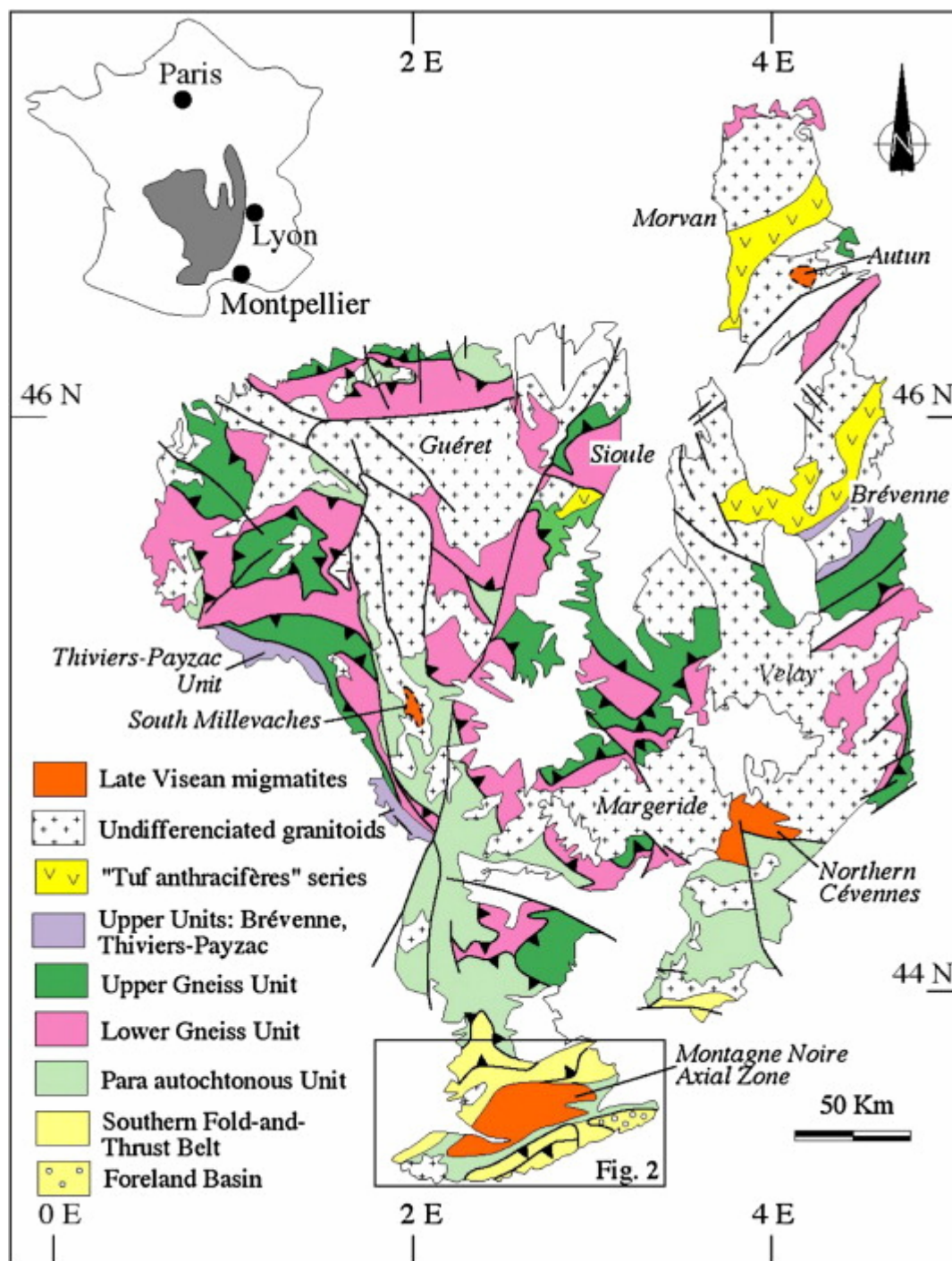


Fig. 1. : Structural map of the Variscan French Massif Central (modified from [Ledru et al., 1989], [Faure et al., 2005], [Faure et al., 2009a] and [Faure et al., 2009b]) showing the location of the Montagne Noire in the southern part of the massif.

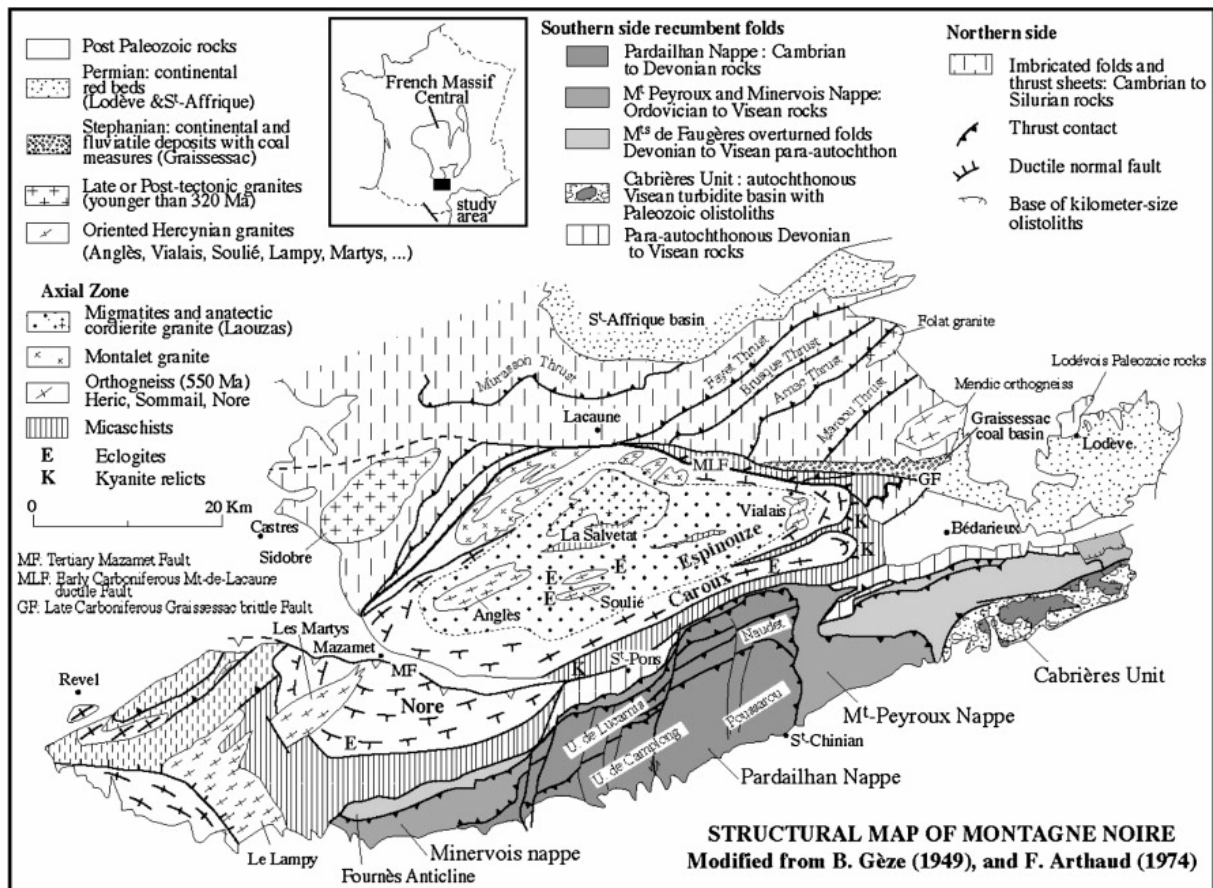


Fig. 2. : Structural map of the Montagne Noire (modified from [Gèze, 1949] and [Arthaud, 1970]). The Montagne Noire is subdivided into a gneissic–migmatitic–granitic axial zone and Paleozoic recumbent folds in the northern and southern flanks.

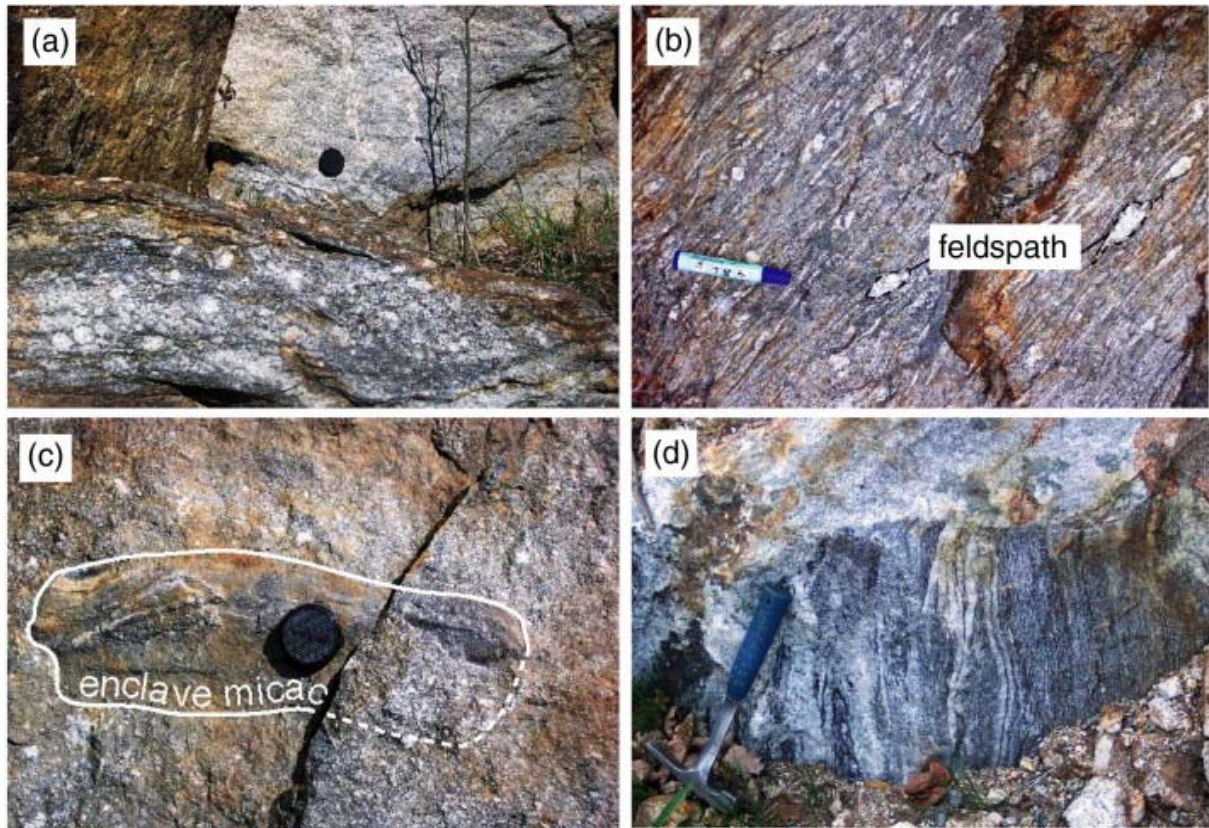


Fig. 3. : Examples of migmatites observed in the Montagne Noire axial zone showing various stages of crustal melting; a: highly melted Ourtigas metatexite; b partially melted Ourtigas orthogneiss with preserved foliation and K-feldspar megacrysts; c: Laouzas anatectic granite with paragneiss enclave oriented in the granite foliation; d: La Salvetat metatexite.



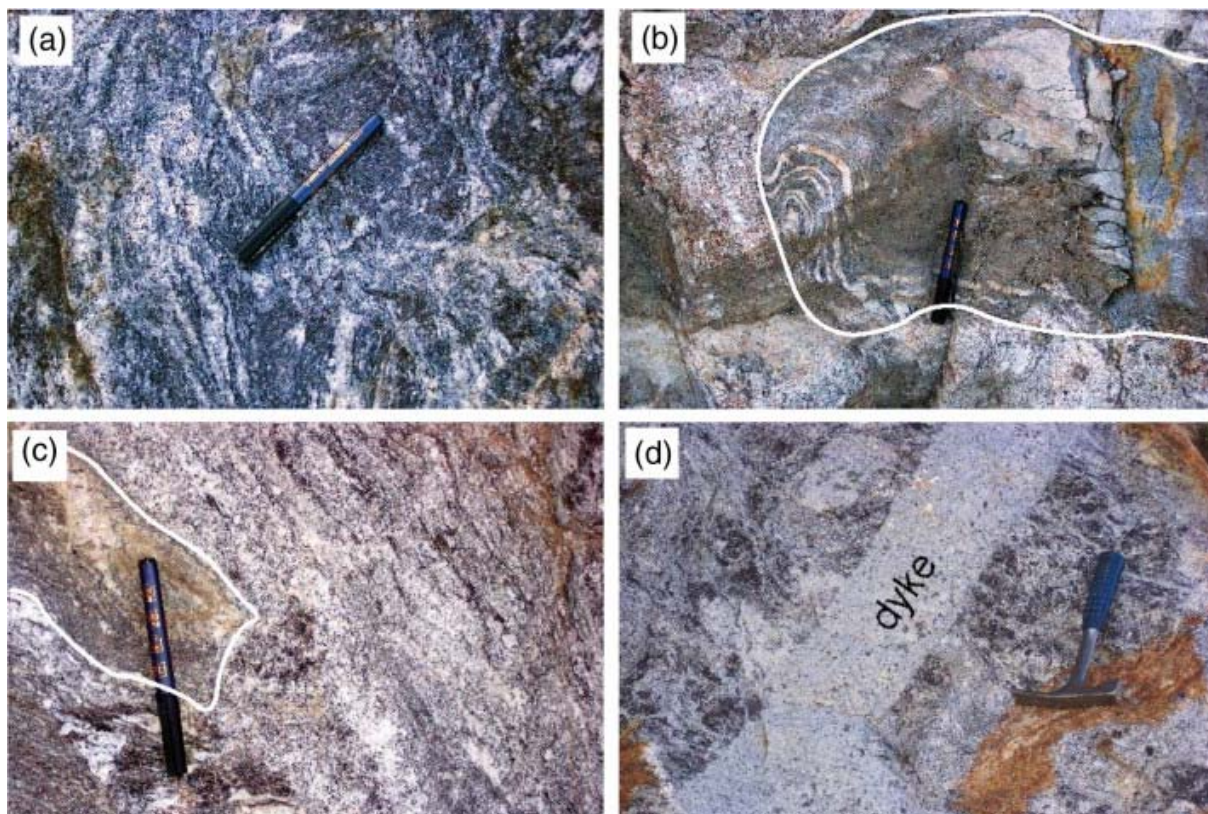


Fig. 4. : Examples of La Salvetat migmatites and anatectic granitoids; a: diatexite; b: Laouzas anatectic granite with metatexite enclave; c: metatexite with paragneiss enclave; d: Laouzas anatectic granitoid cross cut by fine grained aplitic dyke.

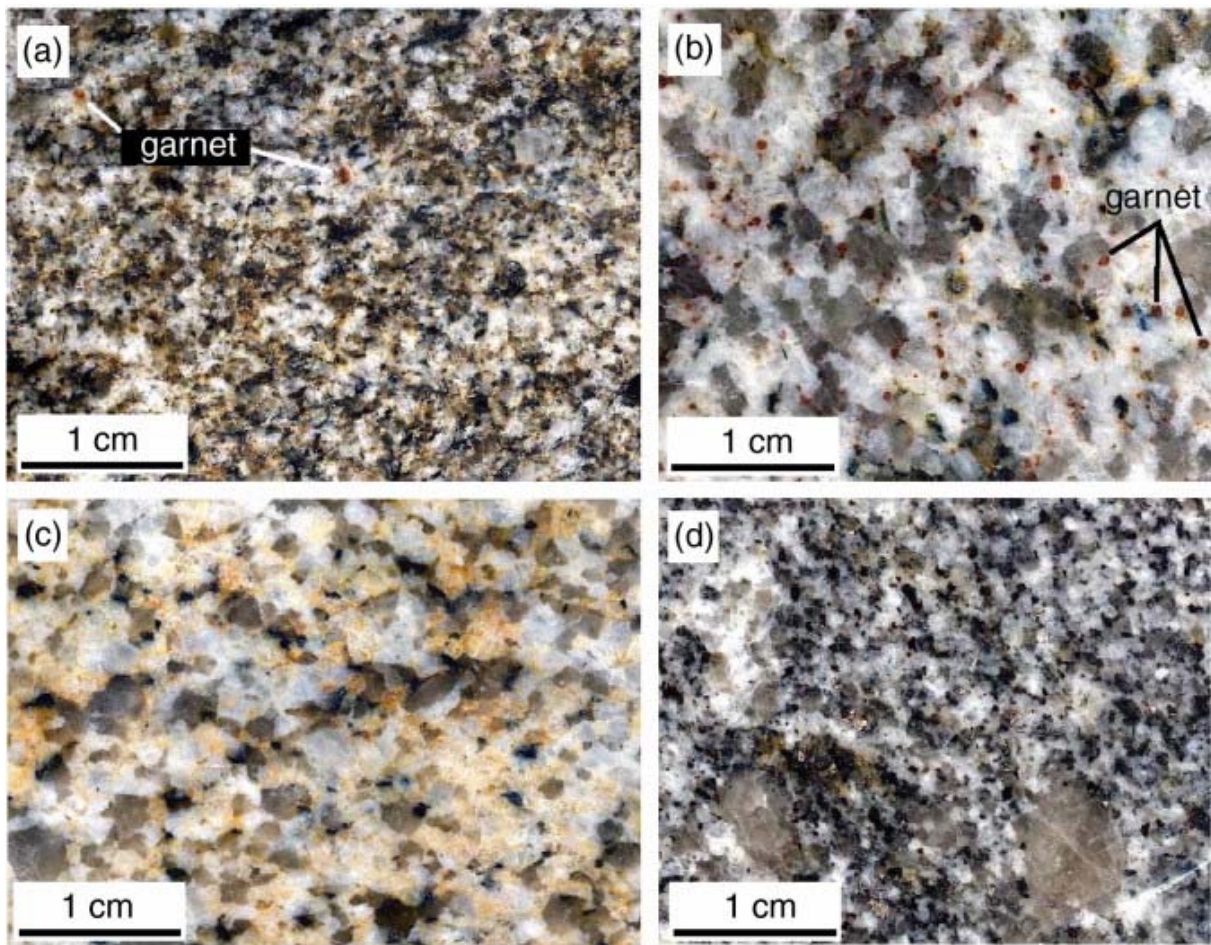


Fig. 5. : Examples of granitoids. a: fine grained Montalet biotite-rich facies with a small amount of garnet; b: medium grained Montalet garnet-rich facies; c: Vialais leucogranite; d: Anglès biotite granite with K-feldspar megacrysts.



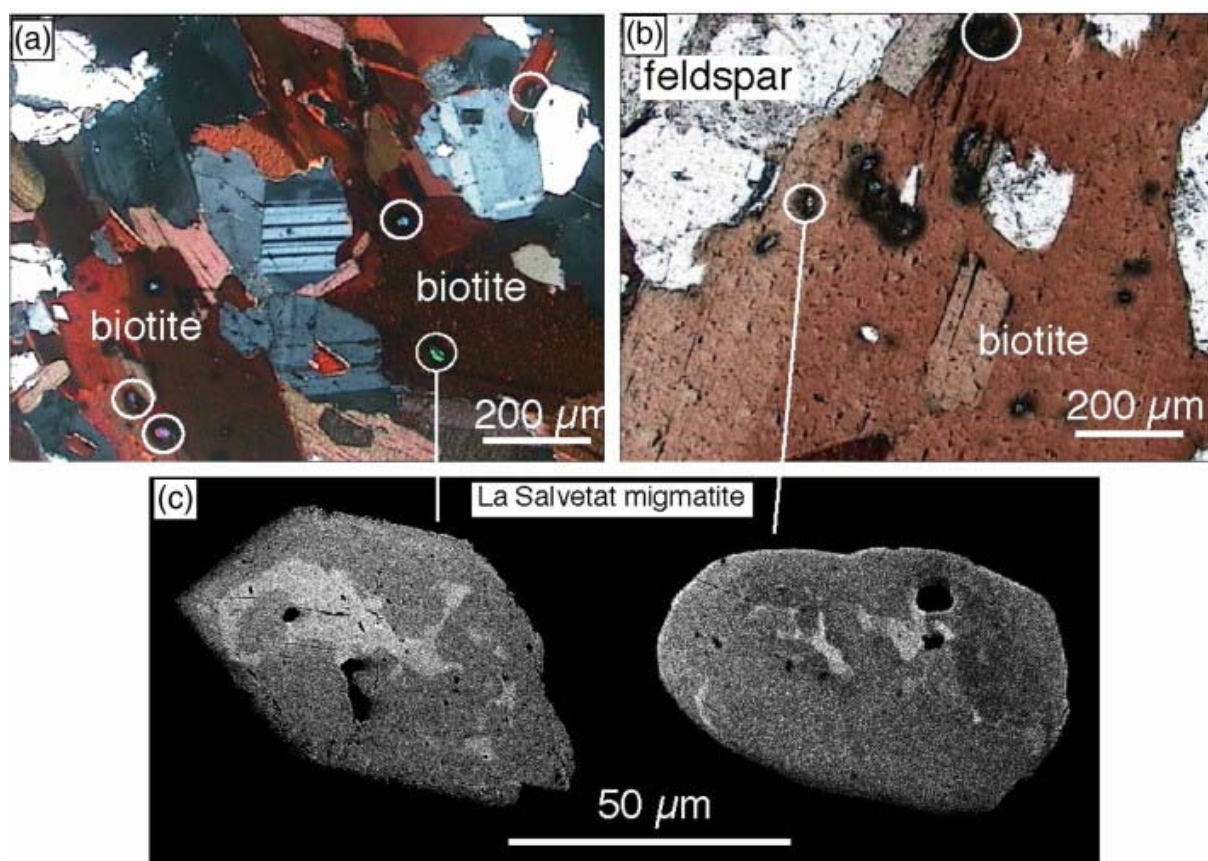


Fig. 6. : Microscopic aspect of La Salvetat migmatite. a: monazite (circled) inclusions in biotite with pleochroic halos, b: detail of monazite inclusions in biotite, c: Back-scattered electron SEM images of monazite grains with subhedral or rounded shapes and patchy zoning, light areas correspond to Th-rich domains.

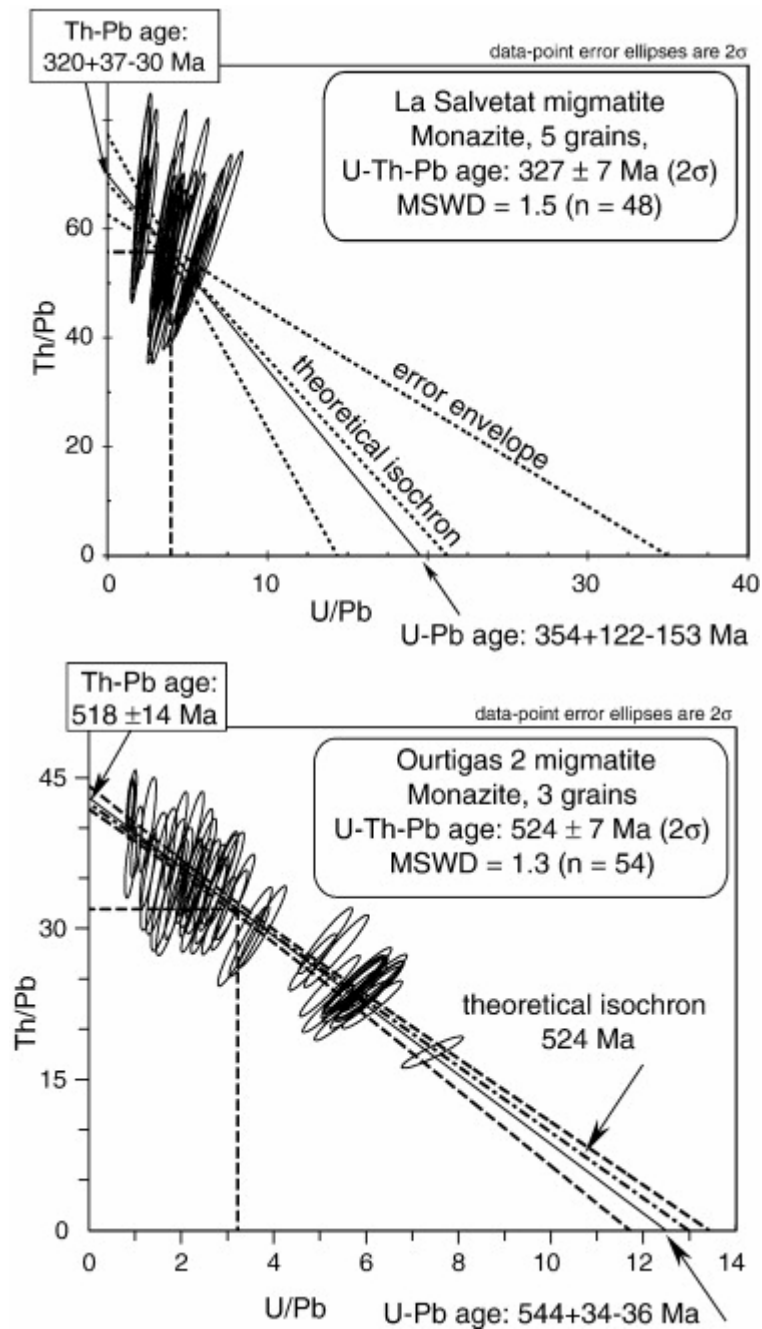


Fig. 7. : U/Th vs. U/Pb isochron diagrams of monazite from La Salvetat and Ourtigas (Ourt 2) migmatites. For each diagram, the regression line is close to the theoretical isochron. The age determined at the intersections of the regression line with the U/Pb and Th/Pb axes support the validity of the calculated age determined at the centroid of the analyses represented by their error ellipse.

Table 1. U–Th–Pb monazite average analyzes.

	<b>Pb (ppm) ±</b>	<b>U (ppm) ±</b>	<b>Th (ppm) ±</b>	<b>Th/U ±</b>	<b>Isochron age ± 2σ Ma</b>	<b><i>n</i></b>
<i>Migmatite</i>						
Ourt1	1945 ± 473	22544 ± 11484	60542 ± 15041	5 ± 4	326 ± 4	29
Ourt1	1403 ± 393	4644 ± 2930	53513 ± 20255	16 ± 10	458 ± 9	30
Ourt2	1639 ± 397	5453 ± 3234	51865 ± 20470	14 ± 11	524 ± 7	54
Espinouse	1980 ± 433	25297 ± 10597	59201 ± 15287	4 ± 3	322 ± 6	36
Salvetat	1023 ± 186	4124 ± 1425	58133 ± 10731	16 ± 7	327 ± 7	48
<i>Granite</i>						
Vialais	2094 ± 425	25148 ± 9335	65922 ± 17006	3 ± 3	320 ± 3	92
Anglès	1106 ± 198	10003 ± 4137	45417 ± 5565	5 ± 2	325 ± 7	56
Montalet- bi	2196 ± 446	9246 ± 7224	68681 ± 10869	19 ± 17	499 ± 6	96
Montalet- gt	937 ± 180	7060 ± 3481	42572 ± 3427	7 ± 3	327 ± 7	58
Laouzas	1133 ± 216	4660 ± 1377	66298 ± 13792	15 ± 5	333 ± 6	54



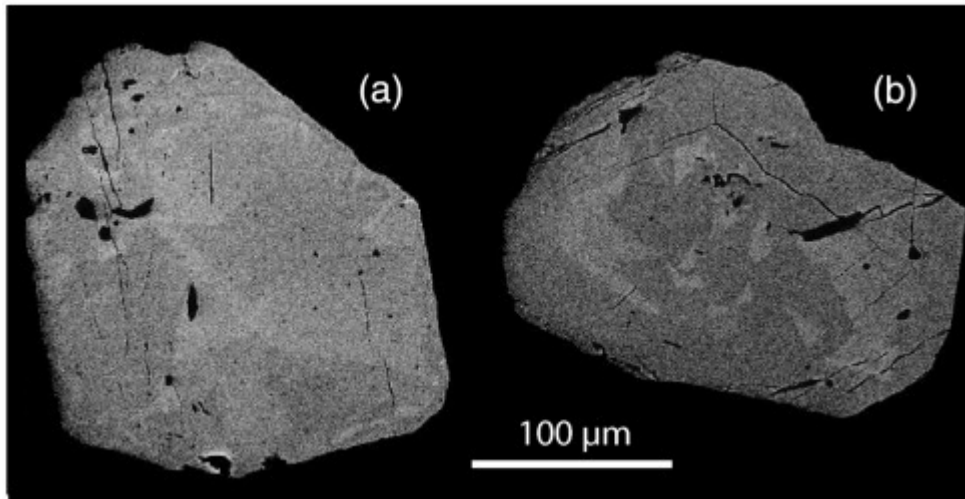


Fig. 8. : Examples of back-scattered electron SEM images of analyzed monazite grains, in the Ourtigas migmatites: a) Ourt 1, b) Ourt 2 samples. The grains exhibit subhedral shapes and patchy zoning, light areas correspond to Th-rich domains.

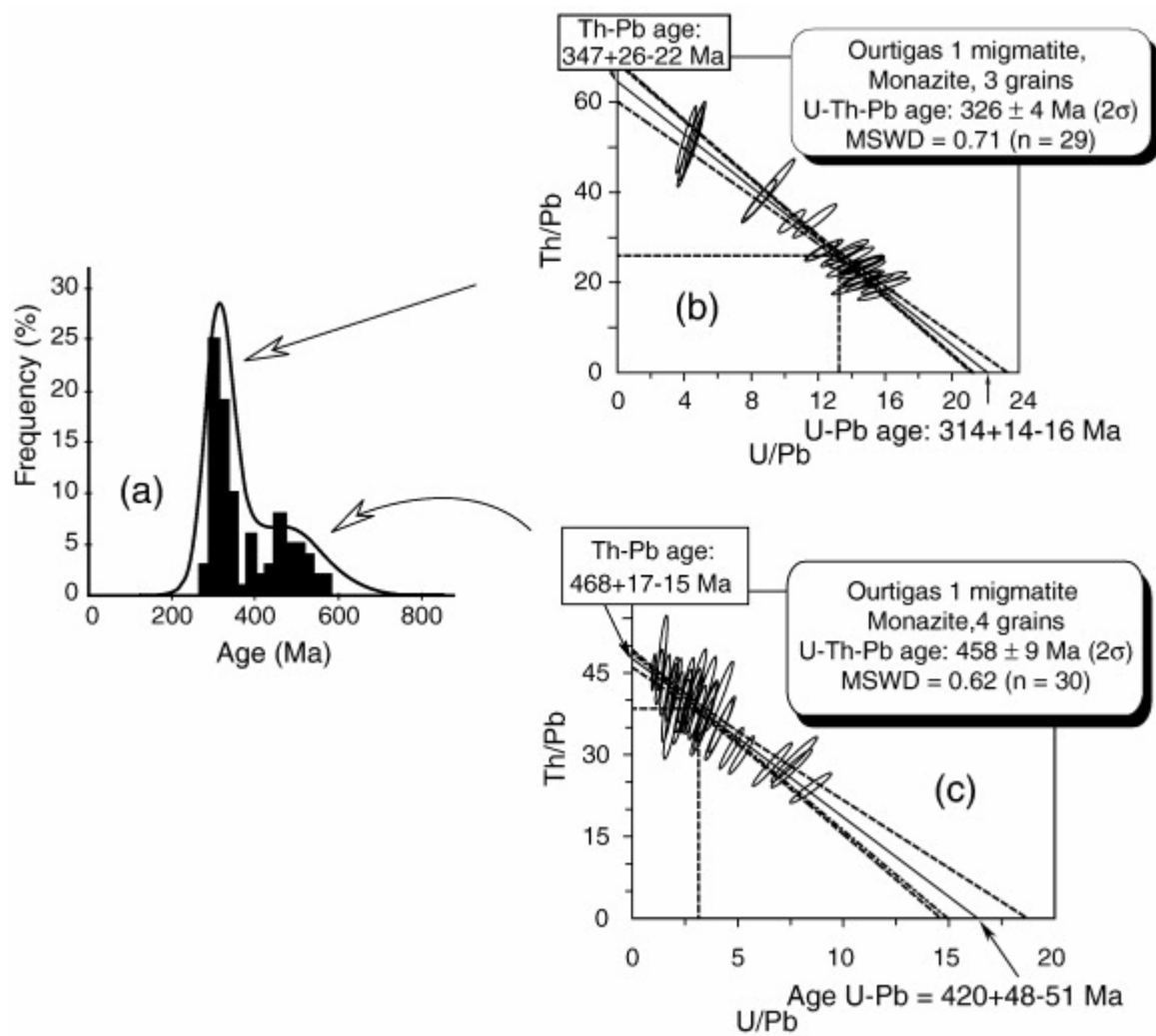


Fig. 9 : Monazite U–Th–Pb dating of the Ourtigas migmatites (Ourt 1). Individual age frequency diagram (a) shows a bimodal distribution. The highest peak corresponds to grains yielding a ca 326 Ma age interpreted as that of partial melting, and the lowest one represents grains yielding a ca 458 Ma age interpreted as that of the orthogneiss protolith (cf. Fig. 3 for outcrop exposure).

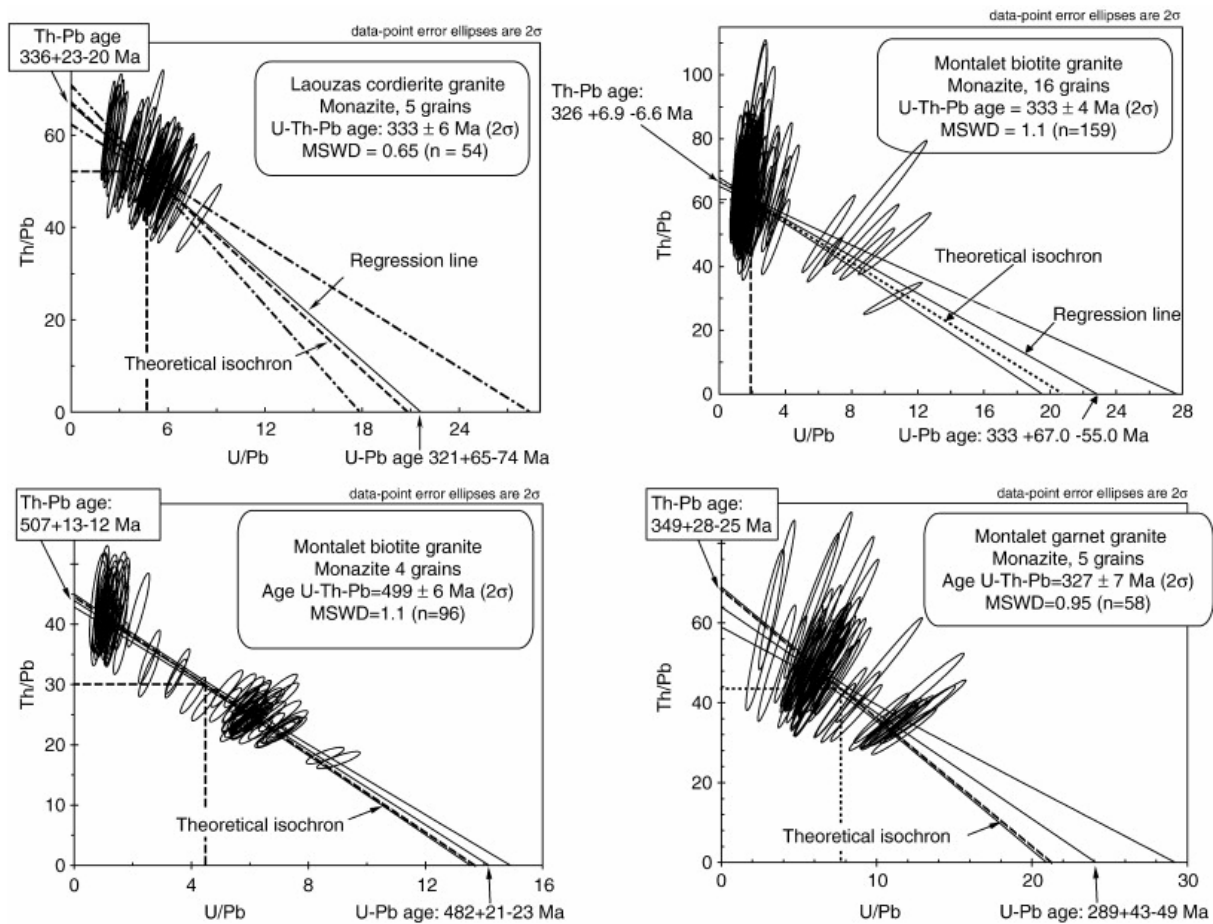


Fig. 10. : U/Th vs. U/Pb isochron diagrams of monazite from Laouzas cordierite granite, and biotite-rich and garnet-rich Montalet granites. For each diagram, the regression line is close to the theoretical isochron. The age determined at the intersections of the regression line with the U/Pb and Th/Pb axes support the validity of the calculated age determined at the centroid of the population. Three ages range around 333–327 Ma, whereas the last age at ca 500 Ma is interpreted as that on inherited grains from the protolith magmatic source.

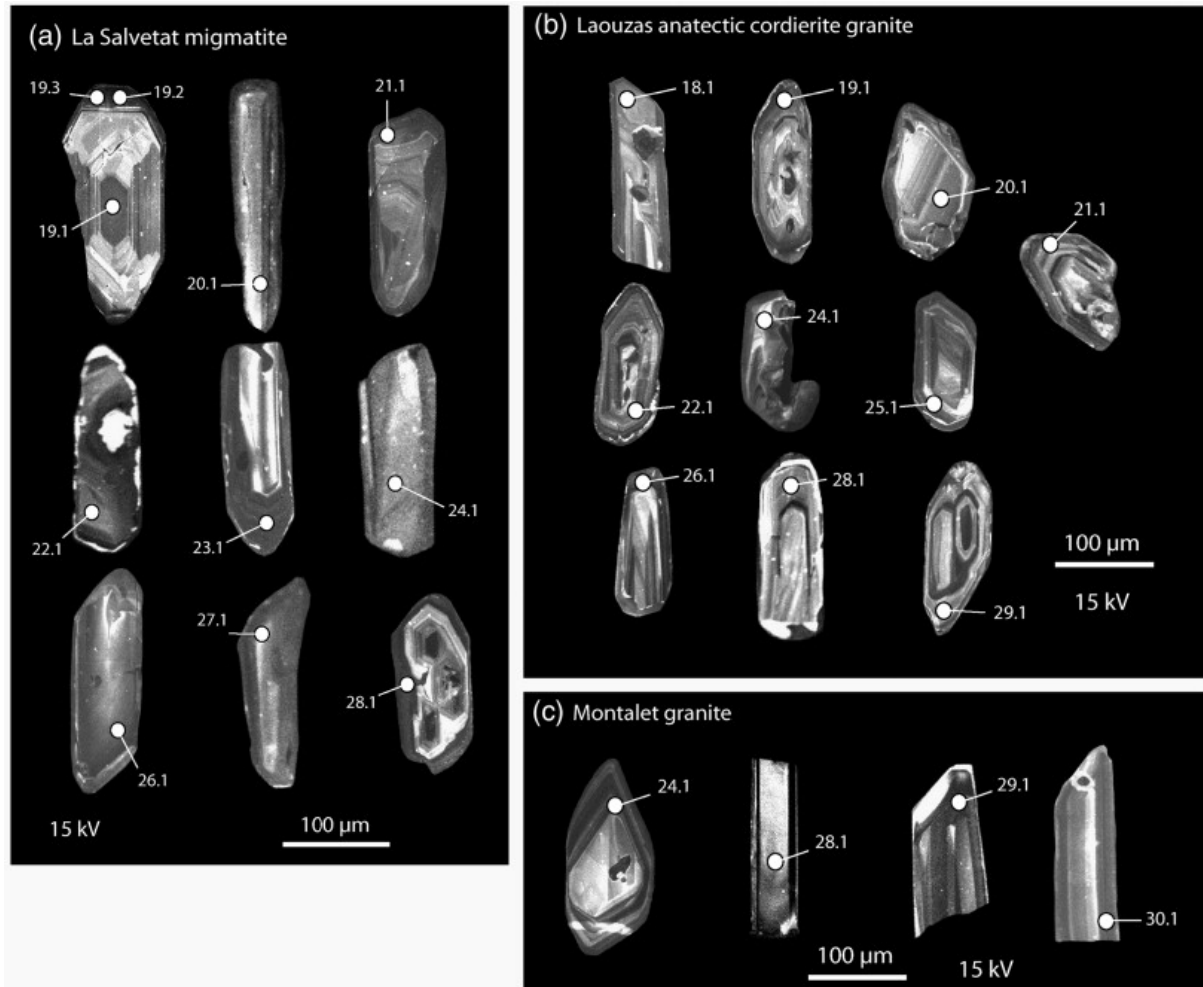


Fig. 11. : Examples of cathodo-luminescence images of the dated zircon. a — La Salvetat migmatites, b — Laouzas cordierite granite, c — Montalet garnet-rich facies. In La Salvetat migmatite and Laouzas granite, most of zircons are euhedral with inherited cores. Numbers for each zircon refer to analytic [Table 2a], [Table 2b], [Table 2c] and [Table 3].

Table 2a. Ion probe results for La Salvetat Migmatite.

							Total				Radiogenic		Age (Ma)	
Grai n spot	U (pp m)	Th (pp m)	Th/ U	<sup>206</sup> Pb * (pp m)	<sup>204</sup> Pb/ <sup>206</sup> Pb	<i>f</i> <sub>206</sub> %	<sup>238</sup> U/ <sup>206</sup> Pb	±	<sup>207</sup> Pb/ <sup>206</sup> Pb	±	<sup>206</sup> Pb/ <sup>238</sup> U	±	<sup>206</sup> Pb/ <sup>238</sup> U	±
1.1	375	41	0.11	27	—	0.20	11.725	0.27 4	0.05957	0.000 51	0.08512	0.001 99	527	1 2
2.1	326	44	0.14	26	—	< 0.0 1	10.906	0.35 2	0.05846	0.000 23	0.09175	0.002 96	566	1 7
2.2	476	73	0.15	36	—	< 0.0 1	11.272	0.18 1	0.05786	0.000 13	0.08879	0.001 43	548	8
4.1	875	91	0.10	64	0.00003	0.14	11.682	0.28 8	0.05911	0.000 41	0.08548	0.002 11	529	1 3
5.1	148	57	0.38	11	—	< 0.0 1	11.187	0.21 0	0.05792	0.000 28	0.08946	0.001 68	552	1 0
6.1	460	193	0.42	36	—	< 0.0 1	10.919	0.31 3	0.05795	0.000 30	0.09169	0.002 63	566	1 6
7.1	314	192	0.61	22	—	0.19	12.149	0.42 6	0.05901	0.000 57	0.08216	0.002 88	509	1 7
7.2	137	32	0.23	11	—	0.08	10.419	0.64 4	0.06028	0.000 81	0.09591	0.005 92	590	3 5
8.1	460	54	0.12	36	—	< 0.0 1	11.062	0.26 1	0.05839	0.000 23	0.09044	0.002 13	558	1 3
9.1	190	155	0.82	15	—	0.02	11.071	0.30 7	0.05891	0.000 28	0.09031	0.002 51	557	1 5
10.1	497	62	0.13	37	—	< 0.0 1	11.488	0.26 3	0.05817	0.000 24	0.08705	0.001 99	538	1 2
11.1	453	47	0.10	33	—	0.08	11.937	0.23 1	0.05839	0.000 24	0.08370	0.001 62	518	1 0
12.1	291	44	0.15	23	—	0.46	10.971	0.46 1	0.06265	0.000 59	0.09073	0.003 81	560	2 2
12.2	595	11	0.02	26	—	< 0.0 1	19.992	0.50 3	0.05243	0.000 17	0.05004	0.001 26	315	8
12.3	149	46	0.31	11	—	< 0.0 1	11.500	0.17 1	0.05810	0.000 35	0.08697	0.001 29	538	8
13.1	318	73	0.23	25	—	< 0.0 1	10.901	0.26 9	0.05889	0.000 27	0.09174	0.002 27	566	1 3
15.1	241	45	0.19	18	—	0.00	11.416	0.23 4	0.05827	0.000 30	0.08760	0.001 79	541	1 1
16.1	515	53	0.10	38	—	0.06	11.499	0.27 7	0.05870	0.000 27	0.08691	0.002 10	537	1 2
17.1	272	41	0.15	21	—	< 0.0 1	11.295	0.21 8	0.05824	0.000 33	0.08856	0.001 71	547	1 0
18.1	280	35	0.13	21	—	0.18	11.494	0.19 8	0.05965	0.000 30	0.08685	0.001 50	537	9

							Total				Radiogenic		Age (Ma)	
Grai n spot	U (pp m)	Th (pp m)	Th/ U	<sup>206</sup> Pb * (pp m)	<sup>204</sup> Pb/ <sup>206</sup> Pb	<i>f</i> <sub>206</sub> %	<sup>238</sup> U/ <sup>206</sup> Pb	±	<sup>207</sup> Pb/ <sup>206</sup> Pb	±	<sup>206</sup> Pb/ <sup>238</sup> U	±	<sup>206</sup> Pb/ <sup>238</sup> U	±
19.1	329	419	1.27	27	—	< 0.0 1	10.629	0.19 3	0.05838	0.000 48	0.09420	0.001 71	580	1 0
19.2	894	12	0.01	38	0.00079	1.34	20.311	0.27 0	0.06332	0.000 66	0.04857	0.000 65	306	4
<i>19.3</i>	<i>508</i>	<i>80</i>	<i>0.16</i>	<i>35</i>	<i>0.00001</i>	<i>0.03</i>	<i>12.395</i>	<i>0.11 0</i>	<i>0.05744</i>	<i>0.000 18</i>	<i>0.08066</i>	<i>0.000 72</i>	<i>500</i>	<i>4</i>
20.1	668	69	0.10	44	0.00024	0.42	12.915	0.16 8	0.06013	0.000 15	0.07710	0.001 00	479	6
<i>21.1</i>	<i>447</i>	<i>35</i>	<i>0.08</i>	<i>38</i>	<i>0.00024</i>	<i>0.07</i>	<i>10.041</i>	<i>0.27 8</i>	<i>0.06084</i>	<i>0.000 71</i>	<i>0.09952</i>	<i>0.002 76</i>	<i>612</i>	<i>1 6</i>
22.1	447	36	0.08	29	—	0.36	13.195	0.15 8	0.05936	0.000 22	0.07552	0.000 90	469	5
23.1	1182	54	0.05	84	0.00002	0.24	12.023	0.13 5	0.05960	0.000 13	0.08297	0.000 93	514	6
24.1	589	35	0.06	43	—	0.08	11.731	0.14 0	0.05859	0.000 16	0.08517	0.001 01	527	6
25.1	736	52	0.07	55	0.00001	0.08	11.575	0.13 8	0.05879	0.000 10	0.08632	0.001 03	534	6
25.2	675	15	0.02	28	0.00066	0.01	20.684	0.30 8	0.05252	0.000 28	0.04834	0.000 72	304	4
25.3	796	73	0.09	63	0.00021	0.17	10.937	0.09 3	0.06027	0.000 27	0.09128	0.000 78	563	5
26.1	442	445	1.01	30	0.00031	0.72	12.596	0.16 4	0.06284	0.000 66	0.07882	0.001 02	489	6
27.1	661	46	0.07	44	—	0.27	12.965	0.16 5	0.05884	0.000 09	0.07692			

Notes : 1. Uncertainties given at the 1 $\sigma$  level (absolute value).

2. Uncertainty in 91500 reference zircon calibration is included in the above uncertainties.

3. *f*<sub>206</sub>% denotes the pourcentage of <sup>206</sup>Pb that is common Pb.

4. Correction for common Pb for the U/Pb data has been made using the measured <sup>238</sup>U/<sup>206</sup>Pb and <sup>207</sup>Pb/<sup>206</sup>Pb ratios following Tera and Wasserburg (1972) as outlined in Williams (1998).

5. Analyses not included in age calculations are shown in italics.

\*Radiogenic Pb due to <sup>238</sup>U disintegration.

Table 2b. Ion probe results for Montalet granite.

							Total				Radiogenic		Age (Ma)	
Grain spot	U (ppm)	Th (ppm)	Th/U	<sup>206</sup> Pb* (ppm)	<sup>204</sup> Pb/ <sup>206</sup> Pb	<i>f</i> <sub>206</sub> %	<sup>238</sup> U/ <sup>206</sup> Pb	±	<sup>207</sup> Pb/ <sup>206</sup> Pb	±	<sup>206</sup> Pb/ <sup>238</sup> U	±	<sup>206</sup> Pb/ <sup>238</sup> U	±
4.1	2043	417	0.20	89	0.00020	1.56	19.354	0.193	0.06545	0.00082	0.05086	0.00051	320	3
6.1	468	31	0.07	28	—	0.10	14.242	0.199	0.05645	0.00022	0.07014	0.00098	437	6
9.1	4251	3073	0.72	188	0.00186	3.58	18.750	0.266	0.08250	0.00321	0.05142	0.00073	323	4
10.1	2356	1318	0.56	105	—	< 0.01	19.322	0.195	0.05271	0.00012	0.05177	0.00052	325	3
11.1	2053	1059	0.52	88	0.00005	0.02	19.935	0.201	0.05287	0.00011	0.05015	0.00050	315	3
12.1	844	361	0.43	35	—	0.07	20.812	0.263	0.05292	0.00009	0.04802	0.00061	302	4
12.3	638	276	0.43	27	—	0.09	20.535	0.245	0.05317	0.00049	0.04865	0.00058	306	4
14.1	1525	556	0.36	66	0.00105	2.01	19.450	0.266	0.06897	0.00474	0.05038	0.00069	317	4
15.1	333	54	0.16	21	—	0.10	13.680	0.144	0.05685	0.00037	0.07303	0.00077	454	5
15.2	468	24	0.05	30	—	0.05	13.212	0.173	0.05688	0.00023	0.07565	0.00099	470	6
16.1	2128	670	0.31	95	0.00108	2.26	18.714	0.195	0.07134	0.00072	0.05223	0.00054	328	3
17.1	433	214	0.49	19	0.00001	7.16	18.324	0.210	0.11087	0.00153	0.05067	0.00058	319	4
18.1	845	32	0.04	36	0.00034	0.20	20.324	0.236	0.05415	0.00034	0.04911	0.00057	309	3
19.1	187	109	0.58	27	—	0.14	6.050	0.078	0.07322	0.00032	0.16505	0.00213	985	12
20.1	1159	19	0.02	115	0.00057	2.19	8.467	0.110	0.08134	0.00381	0.11552	0.00149	705	9
21.2	223	104	0.47	17	—	0.17	11.330	0.156	0.05983	0.00038	0.08811	0.00121	544	7
23.	283	25	0.0	18	0.0006	1.68	13.59	0.1	0.0697	0.00	0.072	0.00	450	5

							Total				Radiogenic		Age (Ma)	
Grain spot	U (ppm)	Th (ppm)	Th/U	<sup>206</sup> Pb* (ppm)	<sup>204</sup> Pb/ <sup>206</sup> Pb	<i>f</i> <sub>206</sub> %	<sup>238</sup> U/ <sup>206</sup> Pb	±	<sup>207</sup> Pb/ <sup>206</sup> Pb	±	<sup>206</sup> Pb/ <sup>238</sup> U	±	<sup>206</sup> Pb/ <sup>238</sup> U	±
1			9		8		3	55	0	045	33	083		
23.2	192	41	0.21	12	—	0.16	13.501	0.250	0.05751	0.00036	0.07395	0.00137	460	8
24.1	653	26	0.04	30	0.00005	0.62	18.519	0.186	0.05822	0.00021	0.05366	0.00054	337	3
25.1	1412	1209	0.86	143	0.00000	< 0.01	8.462	0.071	0.06296	0.00027	0.11823	0.00099	720	6
26.1	2615	94	0.04	118	0.00011	0.29	18.983	0.209	0.05538	0.00041	0.05253	0.00058	330	4
27.1	2095	7	0.00	103	0.00001	0.24	17.557	0.201	0.05556	0.00020	0.05682	0.00065	356	4
28.1	557	167	0.30	23	—	0.27	20.730	0.225	0.05462	0.00053	0.04811	0.00052	303	3
29.1	939	432	0.46	42	—	0.05	19.268	0.176	0.05330	0.00015	0.05188	0.00047	326	3
30.1	424	202	0.48	19	0.00002	0.13	19.428	0.230	0.05391	0.00050	0.05141			

Notes : 1. Uncertainties given at the 1 $\sigma$  level (absolute value).

2. Uncertainty in 91500 reference zircon calibration is included in the above uncertainties.

3. *f*<sub>206</sub>% denotes the pourcentage of <sup>206</sup>Pb that is common Pb.

4. Correction for common Pb for the U/Pb data has been made using the measured <sup>238</sup>U/<sup>206</sup>Pb and <sup>207</sup>Pb/<sup>206</sup>Pb ratios following Tera and Wasserburg (1972) as outlined in Williams (1998).

5. Analyses not included in age calculations are shown in italics.



Table 2c. Ion probe results for Lauzas granite.

							Total				Radiogenic		Age (Ma)	
Grain spot	U (ppm)	Th (ppm)	Th /U	<sup>206</sup> Pb <sup>b*</sup> (ppm)	<sup>204</sup> Pb/ <sup>206</sup> Pb	<i>f</i> <sub>206</sub> %	<sup>238</sup> U/ <sup>206</sup> Pb	±	<sup>207</sup> Pb/ <sup>206</sup> Pb	±	<sup>206</sup> Pb/ <sup>238</sup> U	±	<sup>206</sup> Pb/ <sup>238</sup> U	±
1.1	318	72	0.23	23	—	0.13	11.700	0.266	0.05899	0.00033	0.08536	0.00194	528	12
2.1	493	66	0.13	37	0.000006	0.06	11.324	0.272	0.05893	0.00023	0.08825	0.00212	545	13
3.1	783	6	0.01	32	—	0.04	20.737	0.532	0.05277	0.00022	0.04820	0.00124	303	8
3.3	272	71	0.26	20	—	0.18	11.677	0.231	0.05943	0.00024	0.08549	0.00169	529	10
4.1	505	60	0.12	38	—	0.09	11.272	0.177	0.05924	0.00022	0.08863	0.00139	547	8
5.1	370	77	0.21	28	0.000090	0.14	11.445	0.295	0.05944	0.00044	0.08725	0.00225	539	13
6.1	334	32	0.09	23	—	0.17	12.333	0.286	0.05870	0.00037	0.08094	0.00188	502	11
7.1	177	51	0.29	14	0.000174	0.07	11.242	0.257	0.05906	0.00025	0.08889	0.00203	549	12
8.1	352	72	0.20	26	—	0.10	11.630	0.288	0.05887	0.00031	0.08590	0.00213	531	13
9.1	661	71	0.11	50	—	0.06	11.457	0.300	0.05878	0.00023	0.08723	0.00228	539	14
10.1	237	42	0.18	17	—	0.16	11.809	0.210	0.05913	0.00031	0.08455	0.00150	523	9
12.1	363	80	0.22	28	—	< 0.01	11.325	0.253	0.05826	0.00030	0.08832	0.00198	546	12
13.1	534	4	0.01	22	0.001203	1.95	20.779	0.319	0.06801	0.00167	0.04719	0.00072	297	4
13.2	96	40	0.41	8	0.000084	< 0.01	10.661	0.273	0.05890	0.00043	0.09385	0.00240	578	14
14.1	282	31	0.11	21	—	0.07	11.334	0.303	0.05901	0.00043	0.08817	0.00235	545	14
14.2	448	94	0.21	35	—	< 0.01	10.940	0.167	0.05856	0.00012	0.09145	0.00139	564	8
15.1	166	39	0.23	13	—	< 0.01	10.883	0.234	0.05850	0.00039	0.09194	0.00198	567	12

							Total				Radiogenic		Age (Ma)	
Grain spot	U (ppm)	Th (ppm)	Th/U	<sup>206</sup> Pb* (ppm)	<sup>204</sup> Pb/ <sup>206</sup> Pb	<i>f</i> <sub>206</sub> %	<sup>238</sup> U/ <sup>206</sup> Pb	±	<sup>207</sup> Pb/ <sup>206</sup> Pb	±	<sup>206</sup> Pb/ <sup>238</sup> U	±	<sup>206</sup> Pb/ <sup>238</sup> U	±
16.1	359	47	0.13	27	—	0.03	11.290	0.143	0.05868	0.00029	0.08855	0.00112	547	7
16.2	420	34	0.08	32	0.000080	0.31	11.321	0.187	0.06091	0.00038	0.08806	0.00145	544	9
16.3	447	32	0.07	32	—	0.14	12.086	0.218	0.05868	0.00026	0.08263	0.00149	512	9
17.1	773	90	0.12	58	—	0.12	11.469	0.253	0.05922	0.00020	0.08709	0.00192	538	11
18.1	344	33	0.10	26	—	0.02	11.470	0.137	0.05839	0.00020	0.08717	0.00104	539	6
19.1	651	69	0.11	48	0.000045	0.14	11.622	0.149	0.05924	0.00020	0.08592	0.00110	531	7
20.1	241	163	0.67	19	0.000039	0.38	11.045	0.133	0.06185	0.00022	0.09019	0.00109	557	6
21.1	471	80	0.17	37	0.000062	< 0.01	10.874	0.137	0.05838	0.00018	0.09203	0.00116	568	7
22.1	742	77	0.10	52	0.000182	0.31	12.176	0.175	0.05997	0.00021	0.08187	0.00117	507	7
23.1	531	66	0.12	39	0.000098	0.12	11.567	0.138	0.05909	0.00018	0.08635	0.00103	534	6
24.1	577	74	0.13	41	0.000035	0.11	12.233	0.139	0.05825	0.00024	0.08166	0.00093	506	6
25.1	477	70	0.15	34	0.000141	0.10	11.985	0.139	0.05845	0.00035	0.08335	0.00097	516	6
26.1	1041	67	0.06	76	0.000047	0.16	11.800	0.133	0.05917	0.00011	0.08461	0.00096	524	6
28.1	481	31	0.06	37	—	< 0.01	11.085	0.144	0.05814	0.00031	0.09027	0.00117	557	7
29.1	321	92	0.29	24	—	< 0.01	11.465	0.222	0.05783	0.00051	0.08726			

Notes : 1. Uncertainties given at the 1 $\sigma$  level (absolute value).

2. Uncertainty in 91500 reference zircon calibration is included in the above uncertainties.

3. *f*<sub>206</sub>% denotes the pourcentage of <sup>206</sup>Pb that is common Pb.

4. Correction for common Pb for the U/Pb data has been made using the measured  $^{238}\text{U}/^{206}\text{Pb}$  and  $^{207}\text{Pb}/^{206}\text{Pb}$  ratios following Tera and Wasserburg (1972) as outlined in Williams (1998).

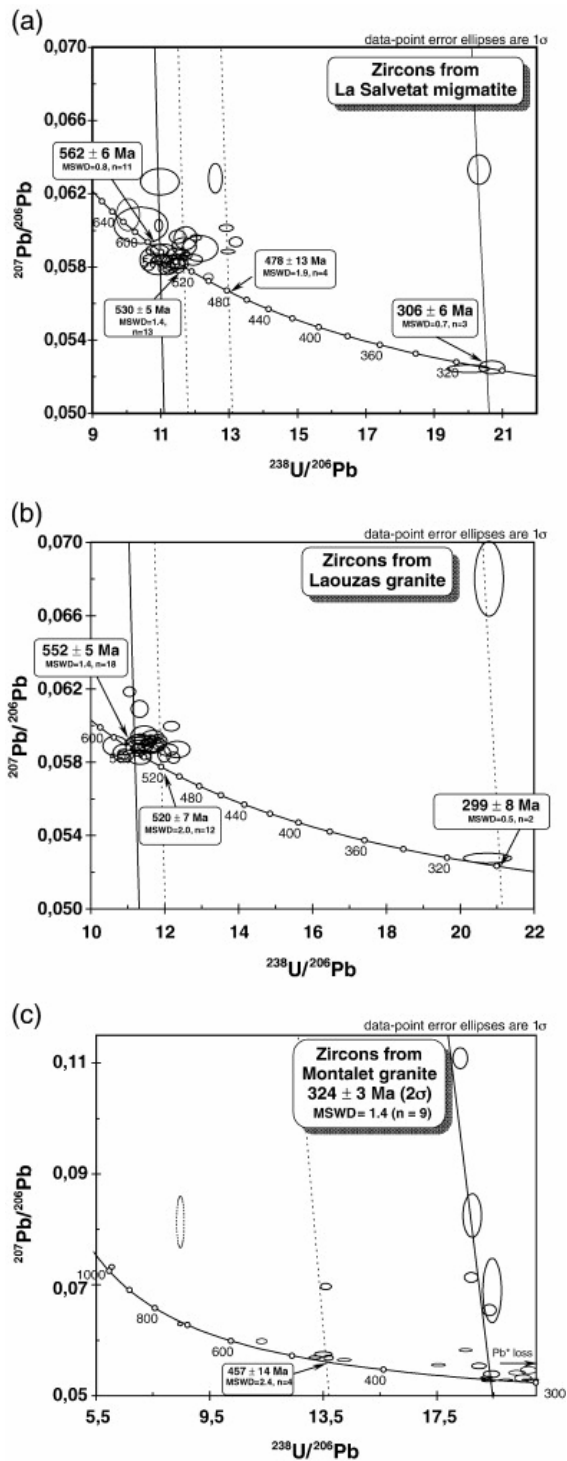


Fig. 12. : Tera–Wasserburg diagrams of zircons U–Pb ages for a) La Salvetat migmatites, b) Laouzas cordierite granite, and c) Montalet biotite-rich facies. For the La Salvetat migmatite, the ca 562 Ma age calculated for 11 analyses represents that of inherited grains from the protolith. The 306 Ma age is poorly constrained by only 3 spot analyses. For the Laouzas cordierite granite, the ca 552 Ma age calculated for 18 analyses represents that of inherited grains from the protolith. The ca 300 Ma age calculated for 2 grains only is geologically

discutable as it is younger than the younger granites (cf Fig. 15) intruding into the Laouzas anatectic granite. For the Montalet granite, 324 Ma age calculated for is interpreted as the granite emplacement age,  $457 \pm 14$  Ma age corresponds to inherited grains from the source of the granitic magma.

Table 3. Inherited zircons in La Salvetat migmatite and Montalet granite.

							Radiogenic Ratios									Age(Ma)						
Grain spot	U (ppm)	Th (ppm)	Th/U	<sup>206</sup> Pb* (ppm)	<sup>204</sup> Pb/ <sup>206</sup> Pb	<i>f</i> <sub>206</sub> %	<sup>206</sup> Pb/ <sup>238</sup> U	±	<sup>207</sup> Pb/ <sup>235</sup> U		<sup>207</sup> Pb/ <sup>206</sup> Pb	±	ρ	<sup>206</sup> Pb/ <sup>238</sup> U	±	<sup>207</sup> Pb/ <sup>235</sup> U	±	<sup>207</sup> Pb/ <sup>206</sup> Pb	±	% Conc		
La Salvetat migmatite																						
28.1	111	32	0.29	19	—	0.00	0.1988	0.0047	4.158	0.124	0.1517	0.0185	0.787	1169	25	1666	24	2365	31	49		
Montalet granite																						
8.1	138	64	0.46	52	—	0.00	0.4507	0.0050	11.007	0.192	0.1771	0.0136	0.628	2398	22	2524	16	2626	22	91		
8.2	94	35	0.38	39	—	0.00	0.4861	0.0088	12.420	0.483	0.1853											

Notes : 1. Uncertainties given at the one  $\sigma$  level.

2.  $f_{206}\%$  denotes the percentage of  $^{206}\text{Pb}$  that is common Pb.

3. Correction for common Pb made using the measured  $^{204}\text{Pb}/^{206}\text{Pb}$  ratio.

4. For % Conc. 100% denotes a concordant analysis.

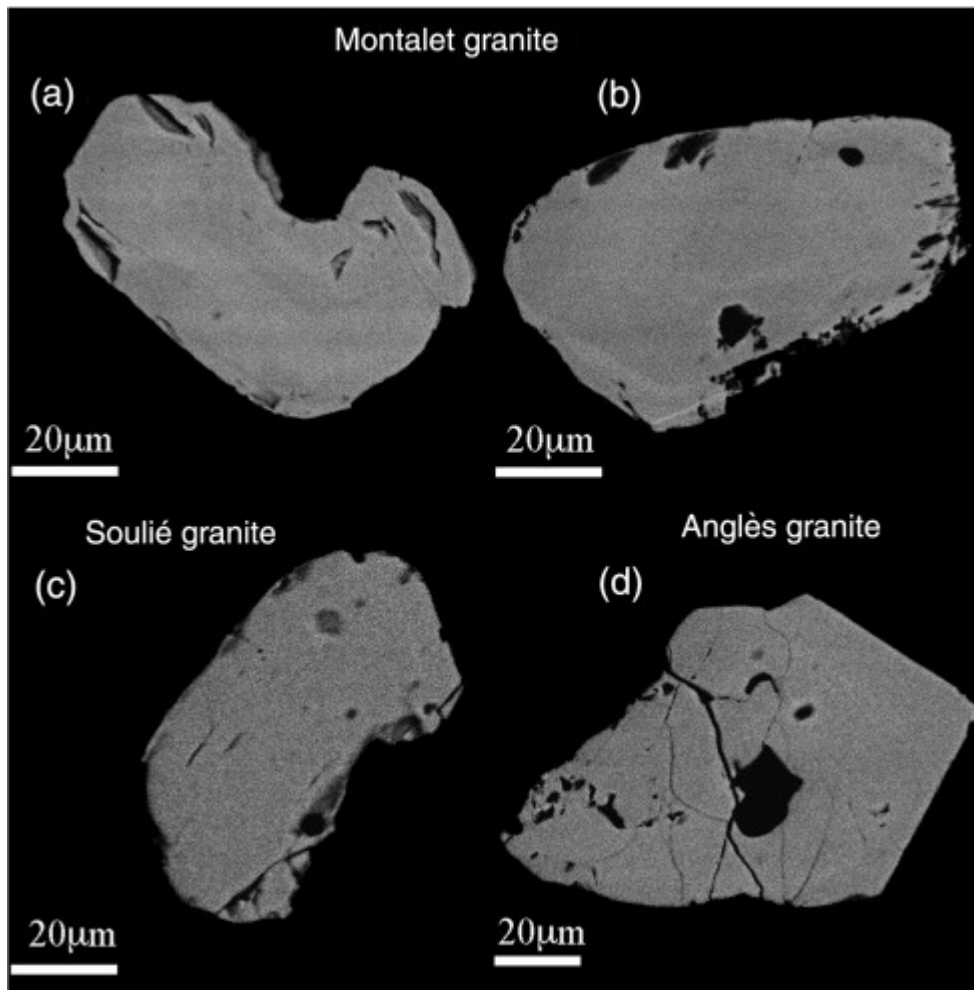


Fig. 13. : Examples of back-scattered electron SEM images of analyzed monazite grains, in the Montalet granite and late migmatitic granites. a) montalet biotite-rich facies, b) Montalet garnet-rich facies, c) Soulié two-micas granite, d) Anglès biotite granite. The grains exhibit subhedral shapes with a faint patchy zoning, and rare inclusions. Light areas correspond to Th-rich domains.

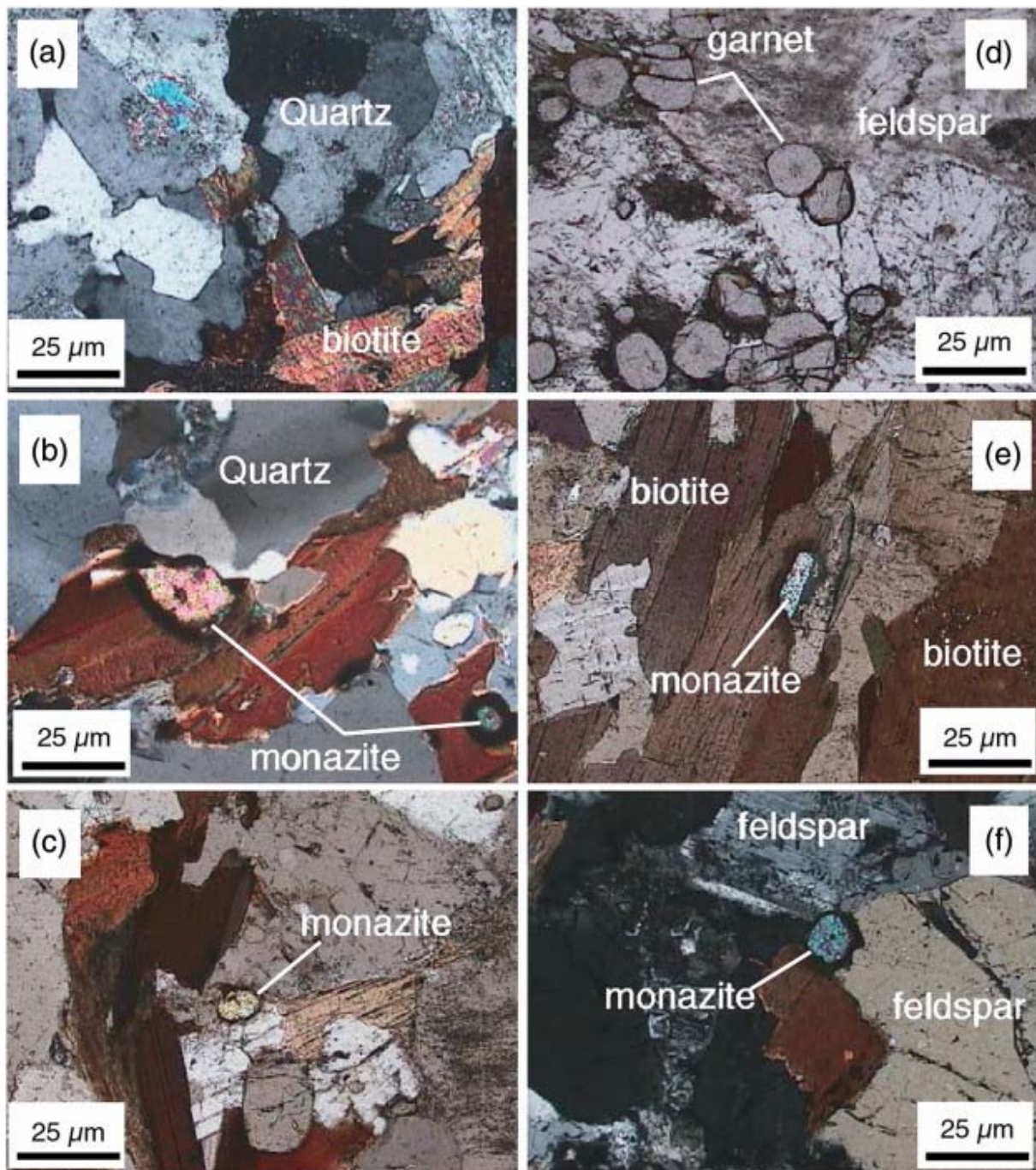


Fig. 14. : Thin sections of Montalet granite showing the textural relationships of monazite with other granite mineral phases. a, b,c: biotite-rich facies; c,d,e: garnet rich facies. Monazite is located either as inclusion in biotite (b, e), or as interstitial phase close to biotite (c) and K-feldspar (d).

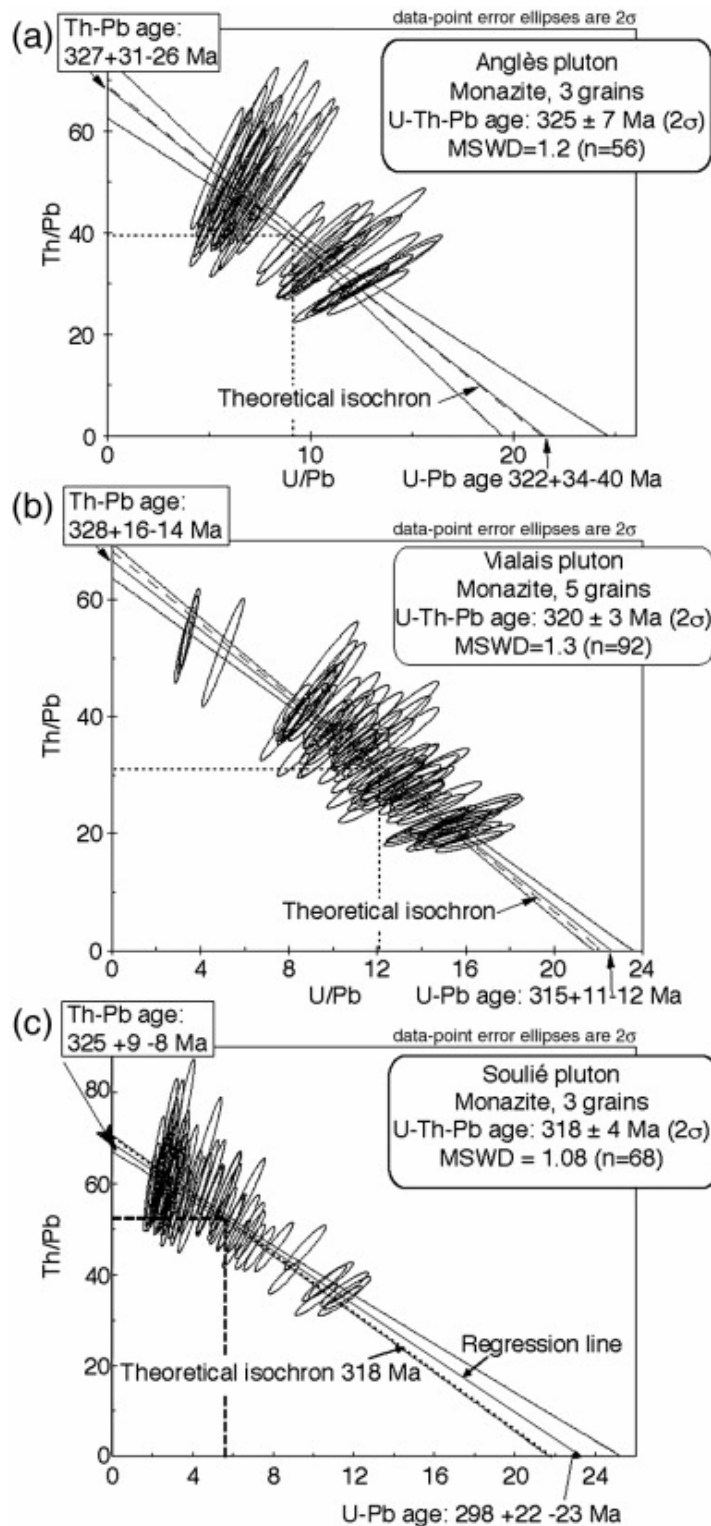


Fig. 15. : U/Th vs. U/Pb isochron diagrams of monazite from late migmatitic granitoids of Anglès (a), Vialais (b), and Soulié (c) plutons. For each diagram, the regression line is close to the theoretical isochron. The age determined at the intersections of the regression line with the U/Pb and Th/Pb axes support the validity of the calculated age determined at the centroid of the population. The three ages, ranging from 325 to 318 Ma, are in good agreement with field observations that show these plutons post-dating the migmatite.



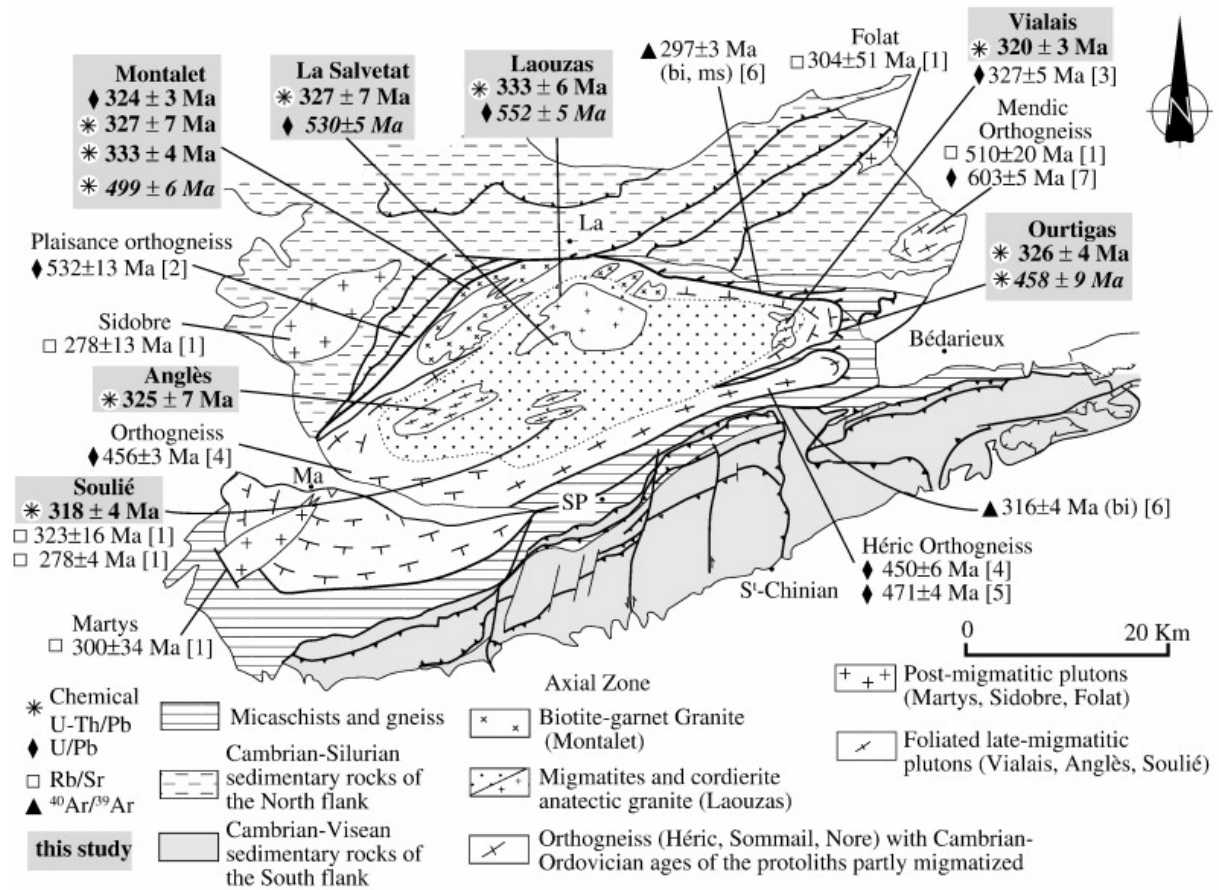


Fig. 16.

Synthetic map of the available radiometric ages for the Axial Zone of the Montagne Noire.

[1]: Hamet and Allègre (1976); [2]: Ducrot et al. (1979); [3]: Matte et al. (1998); [4]: Cocherie et al. (2005b); [5]: Roger et al. (2004); [6]: Maluski et al. (1991); [7]: Lévêque, 1990.



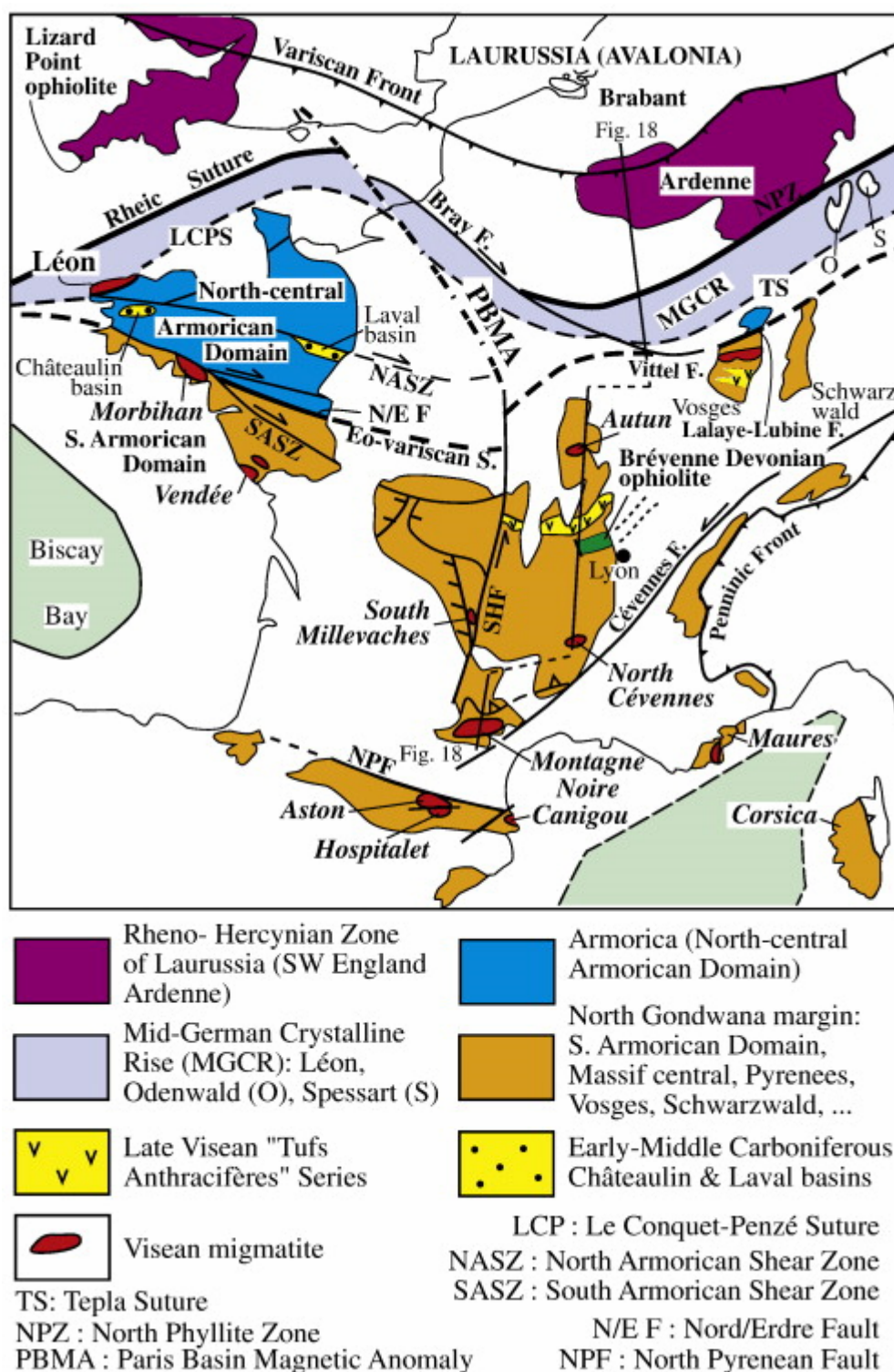


Fig. 17. : Tectonic map of the Variscan French massifs showing the zonation of the belt, and the widespread location of the Middle Carboniferous partial melting, Late Visean–Serpukhovian migmatites, and the Late Visean “Tufs anthracifères” series. MGCR: Mid-german crystalline rise interpreted as a microcontinent extending in the Léon in NW Massif Armoricain. In the MGCR, O: Odenwald, S: Spessart.

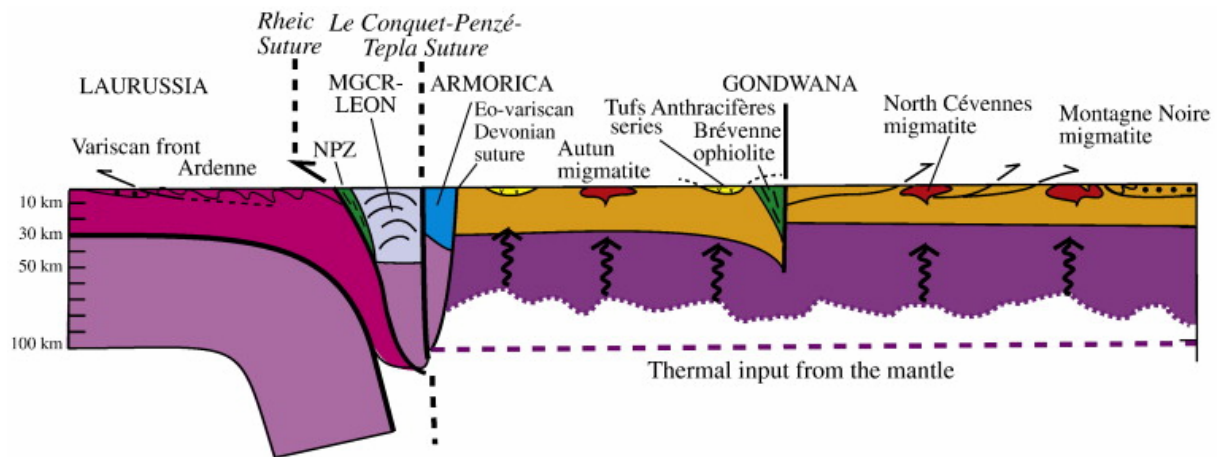


Fig. 18. : Lithosphere-scale interpretative cross-section of the French segment of the Variscan Belt from the Laurussian northern foreland up to the Gondwanian southern foreland during the Late Visean. In order to emphasize the Visean tectono-thermal events, younger granitoids have been omitted. In the French Massif Central, excepting the Eo-variscan suture, other Devonian features are not represented too. After the Early Carboniferous amalgamation of Laurussia, MGCR-Léon, and Gondwana continents was completed by south-directed subduction, the mantle lithosphere was thinned and heat transfer from the asthenosphere triggered the melting of the middle crust giving rise to intracrustal convection and development of the “Tufs anthracifères” series, and anatectic granite–migmatitic complexes throughout the entire belt.